

TP 77-7

Evaluation of the Computation of Wave Direction with Three-Gage Arrays

by Dinorah C. Esteva

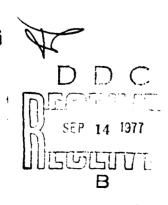
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A description of the collection and analyse array of five pressure sensors near Pt. Mugu, Ca 10 three-gage arrays possible with five gages ar values of the direction of wave propagation. The determination on array orientation relative to i wavelength at the array site is shown by calculationarrow-banded wave trains. Directional results	lifornia, is presented. The e used to compare redundant e dependence of directional ncident wave direction and tions based on simulated

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the maximum accuracy of wave direction determinations with a three-gage array is on the order of $j\pm20^\circ$. This level of accuracy may be expected only for narrow-banded wave trains with periods longer than a lower limit determined at each location by array dimensions and water depth. The field study also indicates narrow-banded wave trains are frequent at this coastal location.

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PREFACE

This report is published to provide guidance to coastal engineers in planning wave data collection in coastal waters for climatology purposes, including wave direction. The popularity of three-gage arrays for proposed wave direction measuring systems makes it necessary to evaluate the capabilities and the limitations of these arrays. The availability of the CERC five-gage array at Pt. Mugu, California, provided a unique opportunity for evaluating the performance of wave recording systems and the directional capabilities of three-gage arrays. The work was carried out under the wave measurement and analyses program of the U.S. Army Coastal Engineering Research Center (CERC).

This report was prepared by Dr. Dinorah C. Esteva under the supervision of Dr. D. Lee Harris, Chief, Coastal Oceanography Branch, Research Division. The author acknowledges the valuable insight and comments provided by Dr. D. Lee Harris and Mr. R. P. Savage, Chief, Research Division, CERC.

Comments on this publication are invited.

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JOHN H. COUSINS

Colonel, Corps of Engineers Commander and Director

CONTENTS

		CONVERSION FACTORS, U.S. CUSTOMARY TO METRIC (SI)	Page 6
	I	INTRODUCTION	7 10 14
	II	FIELD DATA ANALYSIS	14 18 20 28 28 31
		LITERATURE CITED	32
APP	ENDIX A	DERIVATION OF THE EXPRESSION FOR WAVE DIRECTION	35
	В	COMPUTER OUTPUT FOR CROSS-SPECTRA COMPUTATIONS	36
	С	FOURIER COEFFICIENTS FOR A MIXTURE OF THREE SINUSOIDS	48
	D	SPECTRA PLOTS AND COMPUTER OUTPUT FOR SIMULATED OBSERVATIONS	54
	E	HIGH-RESOLUTION SPECTRA FOR FIELD WAVE DATA	79
1	stand	TABLES nt of observations where the largest departure of the dard deviations from the mean in the observations was ndicated (871 observations in 1970)	1 6
2		tional results with 0.01 hertz resolution for two rvations	21
3	Charac	cteristics of simulated wave trains	24
4		tational results at closest spectral frequencies for lated wave trains	26
5	Direct	tional results of high-resolution spectral computations 8-second wave	27
6	ident	ge spread in computed directions for 280 wave trains tified in the high-resolution spectra of 44 field wave rvations	30

CONTENTS--Continued

FIGURES

1	Aerial photo of wave field at Pt. Saint George, California.			Page 8
2	Radar scan of wave field at Nauset, Massachusetts			9
3	Five-gage array dimensions and geometry			11
4	Location of the five-gage array			12
5	Schematics of transducer assembly		•	13
6	Tripod mounting for pressure sensors	•		15
7	Summarized pressure and surface spectra			17
8	Long-crested wave propagating in direction a			19

CONVERSION FACTORS, U.S. CUSTOMARY TO METRIC (SI) UNITS OF MEASUREMENT

U.S. customary units of measurement used in this report can be converted to metric (SI) units as follows:

Multiply	by	To obtain
inches	25.4	millimeters
	2.54	centimeters
square inches	6.452	square centimeters
cubic inches	16.39	cubic centimeters
feet	30.39	centimeters
	0.3048	meters
square feet	0.0929	square meters
cubic feet	0.0283	cubic meters
yards	0.9144	meters
square yards	0.836	square meters
cubic yards	0.7646	cubic meters
miles	1.6093	kilometers
square miles	259.0	hectares
knots	1.8532	kilometers per hour
acres	0.4047	hectares
foot-pounds	1.3558	newton meters
millibars	1.0197×10^{-3}	kilograms per square centimeter
ounces	28.35	grams
pounds	453.6 0.4536	grams kilograms
ton, long	1.0160	metric tons
ton, short	0.9072	metric tons
degrees (angle)	0.1745	radians
Fahrenheit degrees	5/9	Celsius degrees or Kelvins l

¹To obtain Celsius (C) temperature readings from Fahrenheit (F) readings, use formula: C = (5/9) (F - 32).

To obtain Kelvin (K) readings, use formula: K = (5/9) (F - 32) + 273.15.

by Dinorah C. Esteva

I. INTRODUCTION

Wave direction is an important parameter in the solution of many coastal engineering problems. A knowledge of wave direction is essential for (a) estimating the direction and magnitude of sediment transport by waves, (b) using refraction calculations to infer wave conditions at one site from measurements made elsewhere, and (c) verifying theories of wave generation.

Visual observations of wave directions have been collected by shipboard observers for over a century. About 20 years ago the Beach Erosion Board (BEB), predecessor to the Coastal Engineering Research Center (CERC), engaged the assistance of U.S. Coast Guard installations in the collection of visual observations of breaker direction from shore. However, objective determinations of wave direction are desirable without being restricted to location, time of day, or visibility condition. The capability to do so involves the use of wave measuring instruments. Panicker (1971, 1974) presents extensive reviews of reports dealing with the determination of wave direction from instrument records with particular emphasis on those involving sea-surface elevation or pressure records.

In March 1970, CERC installed an array of wave gages at Pt. Mugu, California. Records from the array were to be used to compare redundant values of wave direction and to estimate the level of accuracy of the computations. The available procedures for determining wave direction from an array involved assumptions that had not been thoroughly established. Thus, the records from the array would also be used in a systematic examination of these assumptions, and of the reliability of wave gages.

This study discusses the array performance and the information gained about wave direction. Redundant values of directions were obtained from the 10 three-gage arrays possible with five gages. The mathematical model used assumes that the sea surface is the result of the superposition of a small number of narrow-banded wave trains consisting of long-crested waves traveling in well-defined directions. It was also assumed that only one wave train is present with a particular period. The first assumption is supported by the energy spectra computed at CERC (Thompson, 1974), by aerial photos of the sea surface (Fig. 1), and by radar images of the wave field (Fig. 2). Many published reports include photos similar to that in Figure 1; e.g., McClenan and Harris (1975). Fujinawa (1974, 1975) conjectured that narrow directional spread might be responsible for the incomplete recovery of the true directional spectrum from field records in his computations using high directional resolution.

Average values of wave direction for bands of constant frequency width were computed from cross-spectra between gage pairs. Direction



Figure 1. Aerial photo of wave field at Pt. Saint George, California.

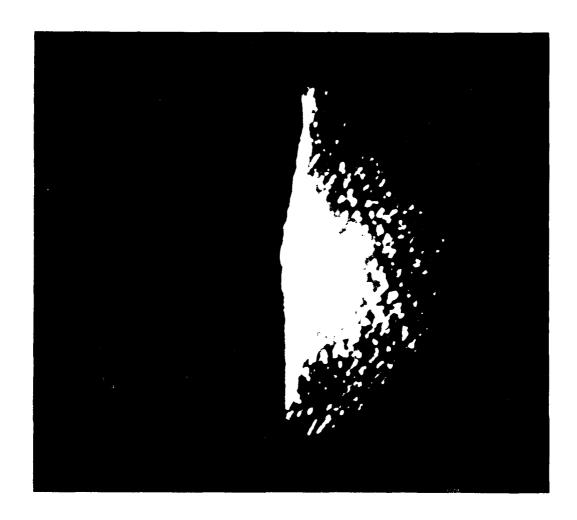


Figure 2. Radar scan of wave field at Nauset, Massachusetts.

estimates for all bands 0.01 herts wide between approximately 30 and 3 seconds were obtained for the 10 arrays. The results displayed discrepancies of the order of 20° for those bands with central periods above 10 seconds and of 180° for those with shorter central periods. It had been expected that the array would yield direction to better than 20° and for periods between 25 and 7 seconds.

To isolate problems associated with the calculations, the propagation of narrow-banded wave trains across the array was simulated in a computer. The computational model was applied to the simulated observations using the maximum frequency resolution available from spectral computations based on 20-minute records. It was found that the directional results obtained with this model are highly dependent on the spectral width, both in frequency and direction, of the wave train involved and on the relationship between wavelength at the site and gage separations. The assigned directions were recovered within 1° for 16-second waves when the frequencies of the spectral components in the wave train differed by 0.003 hertz or more, and the directions were spread within a 5° arc. This frequency separation results in a minimum difference in periods of 0.7 second for waves with periods near 16 seconds.

Application of the same analyses to field wave pressure records with standard deviations above 0.61 meter (2 feet) resulted in an average discrepancy of 20° among computed directions for narrow-banded wave trains with periods longer than 10 seconds. Larger discrepancies resulted for shorter periods. Thus, accuracies no better than 20° can be expected for wave directions resulting from three-gage arrays.

1. The System.

A minimum of three gages is required for a unique determination of wave direction by most proposed models. Since these models make a few assumptions about the nature of ocean waves which have not been established, some redundance was thought to be necessary which would require a minimum of four gages. However, it was agreed that a five-gage array would increase the probability of redundance in the ocean environment. An array was designed at CERC by Leon E. Borgman, statistician-engineer, while on sabbatical leave from the University of California, Berkeley, when the experiment was being planned. He investigated the directional resolving power of several array geometries and concluded that the pattern shown in Figure 5 would be the most suitable for the conditions to be expected at Pt. Mugu (Borgman and Panicker, 1970).

The array was installed off Pt. Mugu, approximately 80.47 kilometers (50 miles) northwest of Los Angeles (Fig. 4), in about 9.14 meters (50 feet) of water, 0.76 meter (2.5 feet) from the bottom. The gages in the array are pressure transducers developed mostly at CERC (Williams, 1969). The heart of the system is a Fairchild pressure transducer which is potted inside a 2-inch Plexiglas tube (Fig. 5) (Peacock, 1974). A plastic tube filled with silicone oil transmits the pressure from the seawater to the pressure transducer. The silicone oil is separated from seawater by

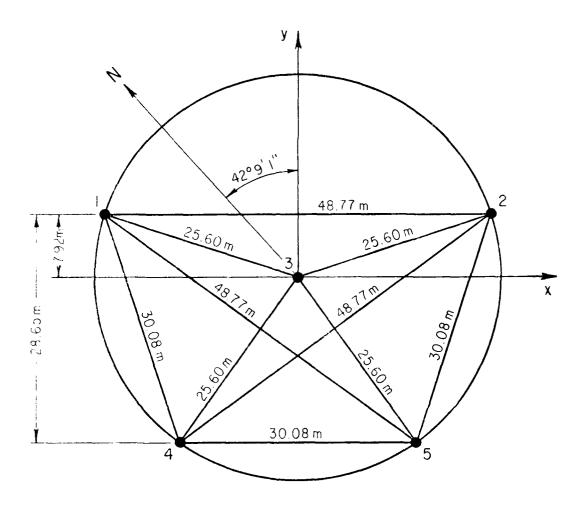


Figure 5. Five-gage array dimensions and geometry.

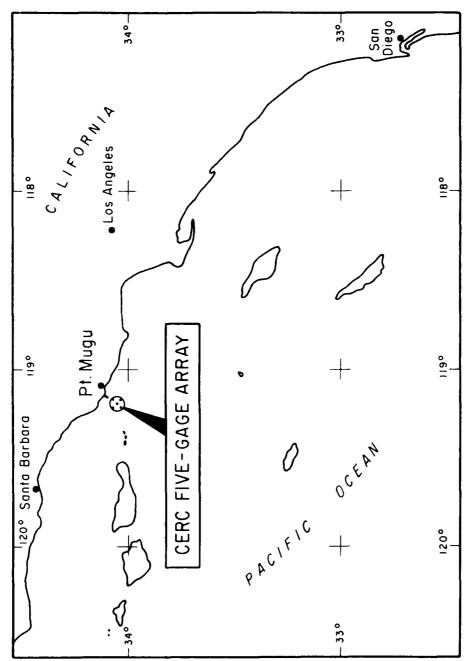


Figure 4. Location of the five-gage array (from Panicker, 1971).

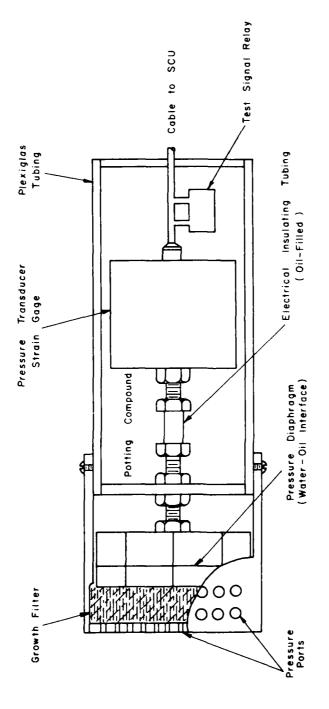


Figure 5. Schematics of transducer assembly (Peacock, 1974).

rubber diaphragms. A Teflon scouring pad dipped in antifouling paint separates the rubber diaphragm from the end of the package which admits the seawater pressure. The instrument is mounted vertically on a tripod (Fig. 6). The signals from the pressure transducers are brought by cables to a recording and transmitting station onshore.

2. Data Collection.

The array went into operation on 27 March 1970. The water pressure at the five gages was registered continuously at a rate of four times a second on digital magnetic tape. Difficulties experienced during the first year with individual sensors were mostly of short duration and were presumed to be due to biological activity. However, major difficulties were experienced with the recording system, and on 16 March 1971 the recorder was disconnected at Pt. Mugu and transferred to the CERC laboratory, then located in Washington, D.C. Recording from all five gages was reinitiated at CERC on 9 April 1971. Records were obtained continuously until 3 January 1972 when the recorder failed. During this period much of the data were useless because of an unacceptable level of high-frequency noise. The source of the noise was difficult to locate and was not eliminated from the signal until shortly before the failure of the recorder in January. Since 2 February 1972, records from three of the five gages at Pt. Mugu have been included in the time-shared recording of waves from east and gulf coast wave stations (Peacock, 1974). In 1972, data were recorded for 20 minutes out of each hour; since February 1973, data have been recorded for 20 minutes out of each 2-hour interval.

II. FIELD DATA ANALYSIS

The five gages in the array provided uninterrupted data for most of the first year of operation. Eight daily observations, each consisting of simultaneous 20-minute records from the five gages, were processed from these data. The observations had starting times within 1.5 hours of the weather synoptic times (0100, 0400, 0700, 1000, 1300, 1600, and 1900 hours P.s.t.). The potential energy in the wave field is proportional to the variance of the time history of sea-surface elevation at a fixed location (Kinsman, 1965). For most conditions, the standard deviation of the surface displacement is one-fourth of the significant wave height. The standard deviation of the pressure at a fixed depth is roughly proportional to the wave height and may also serve as a measure of the wave height.

The standard deviation of the recorded pressure was computed from the records of each of the five gages for eight records each day. The standard deviation of the record from each gage was compared with the average of the five gages. If the standard deviation from any gage differed from the average by more than 20 percent, the record from that gage was deleted and the average recomputed from the remaining gages. A comparison of the individual standard deviations from the mean for that time period is shown in Table 1.



Figure 6. Tripod mounting for pressure sensors.

Table 1. Percent of observations where the largest departure of the standard deviations from the mean in the observations was as indicated (871 observations in 1970).

Deviation from mean (pct)	Percent of Observations
≤ 2	41
3 to 10	52
11 to 20	1
≥ 20	6

This comparison indicates the system operated consistently. The field wave records used for wave direction computation (discussed later) were chosen from the observations in Table 1, for which the standard deviations from all gages differed by 3 percent or less from their mean and for which the average significant wave height (uncompensated for attenuation with depth) was above 0.61 meter.

Fourier analysis provides a reliable procedure for obtaining the periods of the most important waves. A Fourier analysis of sea-surface elevation (or pressure) with time results in the distribution of energy with frequency, usually referred to as the *energy spectrum*. The energy spectrum for the record from gage 5 in each of the eight daily observations was computed using the Fast Fourier Transform algorithm developed by Cooley and Tukey (1967). The first 1,024 seconds of the 20-minute record was used in this computation. Gage 5 was chosen because of a good history of performance.

Fast Fourier Transform computations yield the contribution to the variance at each of a set of frequencies which are harmonics of a fundamental given by the inverse of the record duration, T. In this study, the frequencies of these harmonics are referred to as spectral frequencies, and the corresponding periods, given by their inverse, as spectral periods. The energy content between 32 seconds and 3 seconds was used to normalize the spectrum. The lower limit on period was estimated from the thickness of the water column above the pressure sensors. A summarized spectrum Fig. 7) was formed by combining the energy content in 11 adjacent spectral periods. The band width in the summarized spectrum was slightly arger than 10^{-2} hertz (0.0107 hertz). The energy appearing at each pectral period in the pressure spectrum was compensated for attenuation ith depth by using the classical hydrodynamic pressure correction:

$$F(k,h) = \frac{\cosh kh}{\cosh k\Delta z}, \qquad (1)$$

here k is the wave number, h the water depth, and Δz the vertical istance of sensor from bottom. This resulted in a surface or compensated

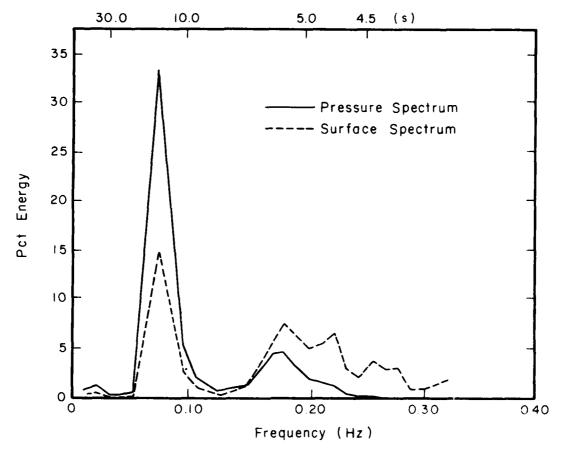


Figure 7. Summarized pressure and surface pectra.

spectrum. Surface root-mean-square (rms) values were obtained from the compensated energy at each spectral period.

Since 1972, four simultaneous 20-minute records from each of the three gages included in the time-shared recording sequence are processed daily. More records are processed during special studies. The records processed from gages 1, 2, and 3 of the Pt. Mugu array start about 0100, 0700, 1300, and 1900 hours (P.s.t.). The significant wave height, the distribution function, the first five moments of the distribution, and the pressure spectra are computed for these records in a study of wave record variability. These records are not analyzed for wave direction.

1. Computation of Wave Direction.

For a long-crested sinusoidal wave with frequency α_m , propagating in direction α_m (Fig. 8), the phase difference between locations 1 and 2 with coordinates (x_1,y_1) and (x_2,y_2) respectively, is given by:

$$\Phi_{12} = k_{\hat{m}} [(x_1 - x_2) \cos \alpha_{\hat{m}} + (y_1 - y_2) \sin \alpha_{\hat{m}}], \qquad (2)$$

where $k_{\hat{m}}^{\perp} = 2\pi/L_{\hat{m}}^{\perp}$ is the wave number associated with frequency $\sigma_{\hat{m}}^{\perp}$, and $L_{\hat{m}}^{\perp}$ is the wavelength. The subscript is used to indicate the possible presence of different wave trains with different frequencies and directions.

The addition of the wave profile, n_3 , at a third noncolinear location allows solving for the sine and cosine of x_m . Thus, a unique solution for the wave direction is obtained from the following equation when the signs of numerator and denominator are considered. (see App. A.):

$$\alpha_{\widehat{m}} = \tan^{-1} \left[\frac{[(x_1 - x_3) \phi_{12} - (x_1 - x_2) \phi_{13}]/D}{[(y_1 - y_2) \phi_{13} - (y_1 - y_3) \phi_{12}]/D} \right], \tag{3}$$

where Φ_{13} is the phase difference between the third and first locations, and D is a function of gage separation.

Phase differences between locations for each different wave period are the only unknowns in the right-hand side of equation (3). Estimates of a representative phase difference between gage pairs for bands of constant frequency width are easily computed from cross-spectra of the wave (pressure) records. These spectra give, for each band, average values of the covariance of the wave records along two perpendicular directions.

Substitution of these representative values of phase difference into equation (3) affords an expedient and economic means of obtaining estimates of a "representative" wave direction for each of the spectral bands, provided the results are of engineering use. The agreement among

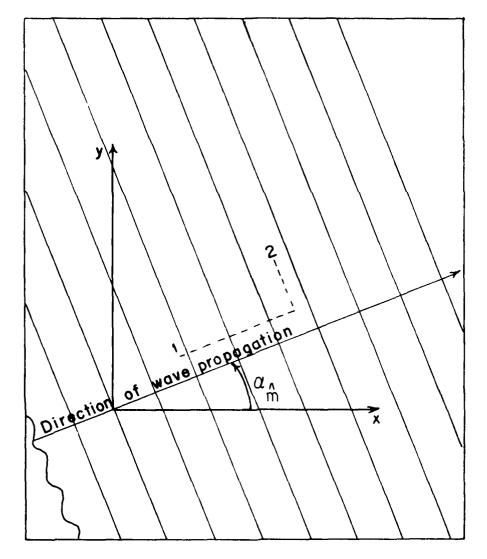


Figure 8. Long-crested wave propagating in direction $\alpha_{\widehat{m}}$.

redundant computations of direction from the 10 three-gage arrays is an indication of the degree of confidence that can be placed on the resulting directions.

Directions were computed for a few observations (using the equation below) for the representative phase difference in each spectral band, where again, the sign of numerator and denominator must be considered.

$$\phi_{12_{\Lambda}} = \tan^{-1} \left[\frac{Quad_{\Lambda}}{Co_{\Lambda}} \right]. \tag{4}$$

The subscript Δ in equation (4) indicates the value is for a given spectral band. The cospectrum, Co, and quadrature spectrum, Quad, are defined as:

$$Co_{\Delta} = \sum_{\Delta} x_1 x_2 \cos \phi_{12}$$

$$Quad_{\Delta} = \sum_{\Delta} x_1 x_2 \sin \phi_{12} , \qquad (5)$$

where $\mathbf{x_1}$ and $\mathbf{x_2}$ are the spectral amplitudes from wave records 1 and 2, and $\frac{\Sigma}{\Lambda}$ indicates summation over the adjacent spectral periods combined to make up a band. Jenkins and Watts (1968) showed that the definitions in equation (5) are equivalent to the more standard definitions based on correlation functions.

A computer output for runs using this approach is presented in Appendix B. Summaries of the results of the observations at 0700 hours, 25 June, and 2010 hours, 28 June, are given in Table 2. The first column in the table gives the period at the center of the band; the second column gives the average percent energy in each band for the five gages. The last 10 columns give the computed "representative" direction of each band for the three-gage array. Directions are measured from the seaward normal with positive values counterclockwise and negative values clockwise. The table shows that in these two observations disagreements in direction of the order of 20° are obtained for the longer period peaks and of 160° for the shorter period peaks. Examination of the computer output in Appendix B indicates the results are typical.

The array had been expected to yield directions to better than 20° for periods between 25 and 7 seconds. Understanding of the problems involved was sought by simulating the propagation through the array of narrow-banded wave trains traveling in a specified direction (discussed in next subsection).

2. Simulated Data Analysis.

In simulating the wave records for use, special consideration was given to wave period and to the difficulties arising in spectral analysis.

Table 2. Directional results with 0.01 hertz resolution for two observations.

	able 2. Dir	ectiona	ii resun	SWILII					o obser	vations.	
Period at center of	Avg. pct energy		1		<u> </u>	nree-ga	ige arra I	ıys			
band	Chergy	1-2-3	1-2-4	1-2-5	1-3-4	1-3-5	1-4-5	2-3-4	2-3-5	2-4-5	3-4-5
(s)						ļ					1
		<u> </u>	25	5 June	1970 a	t 0700	br				
5.99	2.32	173	44	156	50	-42	46	-145	131	-147	-39
6.40	0.46	173	32	157	-44	-44	- 44	147	- 135	- 153	-30
6.87	0.56	64	-23	- 45	- 33	43	- 39	149	- 136	-166	-18
7.42	1.08	167	56	150	-35	50	-42	- 150	- 24	46	-20
8.06	0.75	133	33	162	3	155	49	8	5	-5	0
8.83	1.51	-20	-12	- 9	-19	- 18	0	5	2	- 1	0
9.75	2.01	2	3	2	5	1	13	10	12	16	18
10.89	4.04	5	4	4	4	-2	16	-1	2	8	11
12.34	9.13	21	19	16	16	7	27	14	14	20	21
14.22	27.19	33	30	26	29	16	40	31	28	32	34
16.79	42.92	29	25	22	21	10	33	26	25	27	28
20.48	1.23	1	-1	0	0	- 8	31	26	25	27	28
26.26	0.24	74	119	121	117	131	120	178	144	109	135
			28	3 June	1970 ai	2010	hr				
5.99	7.44	169	52	136	-38	-31	-28	- 148	-136	- 155	-23
6.40	10.85	47	-38	-33	40	-34	- 33	-148	-138	-157	- 25
6.87	16.56	40	- 34	- 29	-36	30	-27	-151	- 22	46	21
7.42	9.75	- 40	33	29	35	-3 0	-27	-152	-20	48	-20
8.06	19.61	40	- 33	- 29	-36	- 30	27	152	-22	49	- 21
8.83	7.65	- 37	- 29	-24	-32	- 27	20	-18	-15	-14	-14
9.75	0.96	16	15	-12	-16	-16	-2	3	-3	-4	0
10.89	1.10	0	0	0	1	-7	25	-3	8	21	24
12.34	2.34	_ 23	13	11	5	-4	26	14	14	14	14
14.22	4.46	12	1.2	10	12	4	24	8	10	16	18
16.79	1.61	19	14	12	9	0	28	15	15	21	21
20.48	1.00	108	46	40	25	2	52	40	49	43	34
26.26	0.70	-91	-89	-75	-89	-67	-95	-92	-66	-56	-72

Assume a sinusoid with frequency given as:

$$\alpha = \frac{2\pi \left(\hat{\mathbf{m}} + \delta\right)}{N\Delta t} , \tag{6}$$

where Δt is the interval of time between samples, $|\cdot|$ is less than or equal to 1/2, and $N\Delta t = T$, the record duration.

Equation (6) provides for assigning frequencies which differ from the spectral frequencies. The contribution to the variance at the spectral frequencies of the sampled record is given by S_m^2 as:

$$S_m^2 = a_m^2 + b_m^2 , (7)$$

where a_{m} and b_{m} are the Fourier coefficients.

Harris (1974) showed that for values of m near \hat{m} (i.e., for spectral frequencies near the frequency of the sinusoid), and for \hat{m} far removed from one and N/2, the approximations below are good estimates to the coefficients.

$$a_{m} \neq \frac{A \sin \pi^{2} \cos \left(\phi - \pi^{2}\right)}{\pi \left(\hat{m} - m + \gamma\right)},$$

$$b_{m} = \frac{A \sin \pi \sin \left(\phi - \pi^{2}\right)}{\pi \left(\hat{m} - m + \phi\right)}.$$
(8)

Slow convergence of the energy toward the spectral period closest to the assigned period is clearly indicated. Thus, the energy is spread over adjacent spectral periods. This spreading, due to the finiteness of the record, is usually referred to as syttheore.

The technique routinely used at CERC to decrease spillover is to apply the cosine bell data window as defined by:

$$\hat{y}_{i} = \frac{1}{2} \left[1 - \cos \frac{2\pi t_{i}}{T} \right] y_{i}, \quad i = 1, \dots, N,$$
 (9)

TO THE SECOND

where y_i are the values in the original record. The Fourier coefficient for the resulting function \hat{y}_i , is given by:

$$\hat{\mathbf{a}}_{m} \doteq \frac{A \sin \pi \delta \cos (\phi - \pi \delta)}{2\pi (\hat{\mathbf{m}} - \mathbf{m} + \delta) \left[(\hat{\mathbf{m}} - \mathbf{m} + \delta)^{2} - 1 \right]},$$

$$\frac{\hat{b}_m}{\hat{b}_m} \doteq \frac{A \sin \pi \delta \sin(\phi - \pi \delta)}{2\pi (\hat{m} - m + \delta) \left[(\hat{m} - m + \delta)^2 - 1 \right]}.$$

Thus, convergence is greatly increased and spillover is effectively reduced to three adjacent spectral periods.

The cosine bell data window was applied to the simulated records; therefore, it was sufficient to consider wave trains consisting of three sinusoids with nearby periods spread over at most six adjacent spectral periods. In general, the periods of waves in the ocean will differ from the spectral periods. Thus, the sinusoids in each simulated observation were specifically assigned periods differing slightly from the spectral periods by use of equation (6). Central periods of about 8 and 16 seconds were chosen to simulate 8- and 16-second swells. Swells with periods in this range are observed on the west coast.

Each simulated wave record consisted of values of the wave profile at the five gage locations computed at 0.25-second intervals for 17.07 minutes (1.14 seconds), to simulate the sampling rate and record duration ou tomarily used at CERC for field data.

The three sinusoids were assigned specific directions and zero initial phase at the origin of coordinates, and were propagated across the array assuming a constant depth of 9.14 meters. Appendix C shows that the Finite bourser Transform gives the correct phases for three sinusoids thus combined, provided the sinusoids are assigned the same direction, nearly the same amplitudes, and the frequency difference of each component to the nearest spectral frequency is the same. A frequency difference of 0.134/1,024 $\frac{1}{7}$ 0.000120 horiz was chosen.

Transcteristics of the first eight simulated observations are given in Table 3. Rough considerations of refraction using linear theory (McClenan, 1975) yield 22° from the normal to the coastline as the maximum possible direction that waves with a lo-second period may have at the lepth of the array. Directions within 21° N. and 21° S. from the normal were chosen for the first four simulated wave trains. Since waves with an 8-second period may approach the coastline at the array site from a much wider arc, directions up to 60° N. and 60° S. were used for the directions of the fifth to eighth simulated trains.

The computed spectra for those eight simulated observations are shown in Appendix D (Figs. D-1 to D-8). In these figures the variance of the record, proportional to the energy, at each spectral period is plottel versus a linear frequency scale. No grouping of the variance at adjacent spectral periods has been made. These ungrouped spectra are referred to as high-resolution spectra. For spectra computed from 1,024-second records, this high resolution is approximately 0.001 hertz.

The program used to compute these spectra is the same program used at CERC for the analysis of field data. Many spectra of field data computed with this program and summarized by grouping 11 adjacent spectral periods are given by Thompson (1974).

The effect of spillover in the spectrum is shown in the figures of Appendix D. Each spectrum resulted from combining only three sinusoids; however, energy contributions appear at from five to nine adjacent spectral periods.

Table 3. Characteristics of simulated wave trains.

Wave	Period)	Amplitude	Direction
train	(s)	(cm)	(0)
ı	16.75	11.43	21 S.
	15.97	15.24	21 S.
	15.25	11.43	21 S.
2	16.75	11.43	21 S.
	15.97	15.24	15 S.
	15.25	11,43	21 N.
3	16.48	11.43	21 S.
	15.97	15.24	21 S.
	15.72	11.43	21 S.
.1	16.48	11.43	21 S.
	15.97	15.24	15 S.
	15.72	11.43	21 N.
5	8.18	11.43	60 N.
	7.99	15.24	60 N.
	7.80	11.43	60 N.
6	8.18	11.43	60 N.
	7.99	15.24	40 N.
	7.80	11.43	60 S.
7	8.11	11.43	60 N.
	7.99	15.24	60 N.
	7.93	11.43	60 N.
8	8.11	11.43	60 N.
	7.99	15.24	40 N.
	7.93	11.43	60 S.

¹ Approximate value,

The periods assigned to the sinusoids giving rise to simulated observations 1, 2, 5, and 6, differed by exactly three spectral periods (see Table 3). The spectra for these four observations (Figs. D-1, D-2, D-5, and D-6) exhibit three maxima separated by two minima. The energy at these two minima may be interpreted as due to spillover since no sinusoids were combined with the corresponding periods.

The high-resolution spectra from field wave observations are discussed later; however, the use of the minima in these spectra to estimate spillover and noise is discussed here. The average energy at the minima between 25 and 7 seconds in the high-resolution spectra of the field observations was used as a measure of spillover and noise. Only spectral periods displaying an energy content at least twice this "background" energy were interpreted as possibly arising from physical wave components in the wave field.

The average directions resulting from the 10 three-gage arrays are given in Table 4 (last column). Only results of computations at the spectral periods closest to the assigned periods are shown. The table shows that for wave trains 1, 2, and 3 the computed directions for the 10 arrays agree with the input directions to within 1° .

The directional results for wave train 4 are correct only for spectral period 62. The main difference between this train and wave train 2 is the narrower spectral width. Wave train 2 gave the correct directions for all spectral periods; wave train 4 did not.

For simulated wave trains 5 to 8, the average directions seem meaning-less. To determine whether these poor results were due to a programing deficiency, another eight sets of simulated records were generated, interchanging periods and directions. The computer output for the simulated observations is in Appendix D. This appendix and Table 5 should be referenced in the following discussion of additional simulated wave trains.

The directional results for the sinusoids with periods clustered around 16 seconds were of the same quality regardless of the assigned direction. However, the directional results for the 8-second sinusoids indicate that the capability to sense the correct direction for these shorter waves depends on the orientation of the three-gage array relative to the direction of propagation of the incoming wave. The resulting directions, which differed by less than 87° from the assigned directions, are given in Table 5. The top of each column in the table shows the shape and orientation of the array. These results are not surprising since the effective gage separation for the different gage pairs varies with orientation relative to direction of wave propagation. Table 5 also shows that the more nearly equilateral arrays have wider direction discernability. The design considerations for the array indicated an effectiveness for wave periods between 25 and 7 seconds (Borgman and Panicker, 1970).

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Table 4. Computational results at closest spectral frequencies for simulated wave trains.

Wave train	Closest spectral frequency	Spectral period	Amplitude	Direction
	(1/1024 hz)	(s)	(cm)	(0)
1	61	16.79	5.62	20 S.
	64	16.00	7.55	20 S.
	67	15.28	5.70	20 S.
2	61	16.79	5.62	21 S.
	64	16.00	7.55	15 S.
	67	15.28	5.70	20 N.
3	62	16.52	5.52	20 S.
	64	16.00	5.35	20 S.
	65	15.75	1.19	20 S.
4	62	16.52	5.52	21 S.
	64	16.00	5.35	25 S.
	65	15.75	1.19	36 N.
5	125	8.19	5.62	142 N.
	128	8.00	7.55	142 N.
	131	7.82	5.70	142 N.
6	125	8.19	5.62	142 N.
	128	8.00	7.55	78 N.
	131	7.82	5.70	162 S.
7	126	8.13	5.52	142 N.
	128	8.00	5.35	142 N.
	129	7.94	1.19	142 N.
8	126	8.13	5.52	142 N.
	128	8.00	7.14	76 N.
	129	7.94	5.83	172 S.

 Table 5. Directional results of high-resolution spectral computations for 8-second wave.

Assigned					Three-gage arrays	ge arrays				
direction¹ (°)	1.2.3	1.24	1-2-5	1.3.4	1-3-5	<u>.</u> ₹∇	1.00 m	2.3.5	1 7	3.4-5
09 N.				59 N.	.N 09	59 N.		.Z. 0.9		
40 N.	40 N.	40 N.	40 N.	40 N.	V 04	Z		Z 2		39 N.
21 N.	21 N.	20 N.	21 N	21 N.	20 N.	20 N.	-	2 Z		S S S
15 N.	1. Z 4.	14 N.	7 Y	Z Z	4 Z	Z.		z t		ź c
15 S.	14 S.	14 S.		14 S.			14 S.	14 S.	14 S.	15 S.
21 S.	21 S.	20 S.		20 S.			20 S.	20 S.	20 S.	24 S.
40 S.	40 S.	39 S.		39 S.			39 S.	40 S.	39.8	39 S.
60 S.				60 S.			60 S.	59 S.		60 S.

1 Results differing by 87° or more from the assigned directions have been omitted.

The consistent and exact recovery of assigned directions achieved for the simulated 16-record wave trains is in part due to the use of the high computational resolution available in a large computer. Use of less exact data as available from recording instruments is expected to result in a less consistent and accurate recovery of the true direction. Simulated observations 1 to 4 were rerun truncating the computed profile values to three digits as is commonly available from recording systems. The effect of this truncation is estimated to have introduced an error of the order of 10.127 centimeter (0.05 inch) in the instantaneous values of the profiles. No appreciable differences in computed directions resulted by this truncation.

3. Identification of Wave Trains from the High-Resolution Spectrum.

A wave train in the real ocean is conceivably made up of several wave components with nearby periods propagating in approximately the same direction. Simplistic idealizations of such wave trains are exemplified by simulated observations 1, 3, 5, and 7. For wave trains 1 and 3, the long-crested wave model gave the correct direction of wave propagation for all 10 combinations (within 5°), not only at those spectral periods closest to the periods of the sinusoids combined, but at several adjacent ones on either side of these periods (see App. D, Figs. D-1 and D-3). At some of these adjacent spectral periods, the contribution to the energy was several times the background level and nearly the same at the five gage locations. Thus, it can be assumed that a wave train in the ocean will give rise to a number of adjacent spectral periods in the highresolution energy spectrum with energy content several times the background level. This background level can be estimated by inspection of the minima in the energy spectrum, as discussed previously. A criterion for what energy level will be considered "high-energy content" can be set, and groups of adjacent spectral periods in the spectrum with highenergy content identified. These groups may each be assumed to arise from the presence of a wave train in the field with a mean wave period within the range of spectral periods in the group (a wave packet). The number of adjacent spectral periods in each group will be used as a measure of the width of the energy peak in the spectrum and indicates the spread in periods of the wave train. The spread in computed directions at adjacent spectral periods in a group is an indication of the degree of directional organization in the wave train. Large spread in directional results may indicate the possibility that crossing wave trains with nearly the same period are present. As results for simulated wave train 4 indicate, the long-crested model based on the assumption of a single wave train at each frequency is not suitable for a determination of wave direction in such cases. Multiple wave trains at a single frequency may result from refraction around a shoal or island or from reflection by a vertical wall.

4. Spectra and Direction of Wave Propagation for Field Data.

The energy and direction of wave propagation at each spectral period were computed for 44 field observations where the average uncompensated

significant wave height was over 2 feet and the discrepancy of individual standard deviations from their mean was 3 percent or less. Plots of the high-resolution spectra for these observations are in Appendix E. The vertical lines represent the energy contribution at each spectral period. The background level for each observation was estimated from the minima between 25 and 7 seconds in the spectra. Spectral periods in this range with energy content above twice the estimated background energy were identified. Contributions to the energy satisfying this criterion at adjacent spectral periods were considered as arising from the same wave train. The number of adjacent periods in each train was used as a measure of the spectral width of the train. The energy had to be above the chosen level at all five gages for the spectral period to be included in the group. The spectral period among these showing maximum energy was taken as the "period" of the wave train.

Directions were computed at all the spectral periods in each train for the 10 arrays. The total spread among these directions was found, and an average total spread was computed for the trains having the same spectral width. The same was done for the computed directional spread at the period of the train.

Twenty-five percent of the identified wave trains had total directional (computed) spreads above 100° and were not considered further. For 89 percent of the discarded trains, the period of the train was under 9.4 seconds. Thus, all trains with periods under 9.4 seconds were discarded.

Results for the different spectral widths for the trains retained (280) are shown in Table 6. The second column in the table gives the average total directional spread for the corresponding spectral width; the third column gives the average directional spread at the period of the wave train. The last column gives the number of wave trains having the spectral width in the first column.

These results indicate that the total directional spread increased with frequency width. Narrow peaks consisting of from one to three spectral periods are most frequent, and the spread in the direction at the period of the train remains relatively constant. Since the average spread for narrow-banded trains (width ≤ 0.003 hertz) is 21.8°, it is expected that three-gage arrays cannot yield directional results to any better accuracy. The mathematical exercise in Appendix C shows that array dimension is a limiting factor as to the shortest period for which some directional discrimination may be expected. An important factor in the validity of the directional result is the spectral structure of the wave train involved. Only in very special circumstances will the quantities involved in equation C-10 (App. C) combine to give better results.

There are various possible explanations for the large spreads observed in the directional results from field records. For the long-crested wave model to be strictly applicable, it is important that:

Table 6. Average spread in computed directions for 280 ways trains identified in the high-resolution spectra of 44 field wave observations.

Spectral width	Avg. total spread in direction for all periods	Avg. spread in direction for the first of the train	Cases
hz			(No.
0.001	21.8	21.8	96
0.002	30.9	22.9	58
0.003	33.6	20.2	41
0.004	41.5	20.4	22
0.005	38.7	19.8	21
0.006	43.8	19.8	16
0.007	48.2	17.0	5
0.008	84.5	17.6	8
0.009	53.0	15.3	6
0.010	49.3	13.7	3
0.011	38.0	18.0	1
0.012	81.0	15.0]]
0.013	48.5	15.2	2

- (a) The phase differences be known accurately or that the probable error in their computed values be known.
- (b) The wave crests over the array site be long and straight; thus, the waves must not have undergone appreciable refraction.
- (c) The sea surface be stationary in time for the duration of the record and in space over the span of the array.

The mathematical exercise in Appendix D indicates that the analysis yields accurate phase differences only for strictly monochromatic conditions. When this is not the case, no accurate estimate of the error involved in the computation of direction can be given. This inability is inherent to the computational procedure and cannot be resolved.

Waves with periods over 8 seconds have been and are undergoing refraction at the site of the array. Therefore, the wave crests are not exactly straight. For the longer waves, with wavelengths at the array site several times the gage separations, the curvature will not

introduce much error. This will not be the case for the shorter waves and orientation of the array becomes important.

Because of refraction, the curvature of a wave train changes, perhaps only slightly, while propagating over the array. This change will introduce differences in the direction at each gage and differences in the direction sensed by different gage pairs, causing undetermined additional errors in the computation of direction. To determine the magnitude of these errors, two additional sets of simulated wave records were generated. The periods of the sinusoids combined were those for simulated observation 3; the directions assigned were spread within a 10° arc for the first set and 20° for the second. The last two computer outputs in Appendix D show that spreads of the order of 16° and 32°, respectively, resulted in computed directions.

A stationary condition in time is usually assumed when developing wave directional models. Indications are that this is not strictly applicable at all times.

The three factors discussed above are sufficient to account for the inaccuracies encountered in the computations.

5. Conclusions.

The results of directional computations, for both simulated and field wave data records, indicate three-gage arrays have some capabilities to determine wave direction under certain conditions. These capabilities depend on:

- (a) The dimension of the array and the water depth at the site which place a lower limit on the wave period for which possibly accurate directions may be computed.
 - (b) The orientation of the array for the shorter periods.
- (c) The nature of the wave field; directional results for wave trains with a narrow frequency distribution or where the computed directions differ little at the adjacent spectral periods might be meaningful.

For wave trains with narrow frequency and directional width, and period above 10 seconds, the three-gage arrays at Pt. Mugu yield directions to an estimated accuracy of 20°.

At the Pt. Mugu site, 16-second waves may approach the coastline at angles of 22° or less from the normal. The directional information provided by the array adds little to this and seems hardly cost effective.

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APPENDIX A

DERIVATION OF THE EXPRESSION FOR WAVE DIRECTION

Let the coordinates of nearby gage sites be (x_i, y_i) , i = 1, N; with N = number of gages. The water surface displacement at each site due to the passage of a sinusoidal wave of frequency σ , and amplitude A, traveling in direction α $(\alpha_{\hat{m}}$ in Fig. 8), is given by:

$$\eta_{i} = A \cos\{k[x_{i} \cos \alpha + y_{i} \sin \alpha] - 2\pi\sigma t - \phi\}, \qquad (A-1)$$

where $k=2\pi/L$ is the wave number, L the wavelength, and ϕ the initial phase at the origin.

The phase difference, ϕ_{ij} , between locations i and j for the sinusoid considered is:

$$\Phi_{i,j} = k[(x_i - x_j) \cos \alpha + (y_i - y_j) \sin \alpha].$$
 (A-2)

Thus, for three noncolinear locations,

$$\Phi_{12} = k[(x_1 - x_2) \cos \alpha + (y_1 - y_2) \sin \alpha],$$

$$\Phi_{13} = k[(x_1 - x_3) \cos \alpha + (y_1 - y_3) \sin \alpha], \quad (A-3)$$

suffice for a unique solution of the direction α . Eliminating first the $\sin \alpha$ terms and then the $\cos \alpha$ terms, to obtain:

$$\sin \alpha = \frac{(x_1 - x_3) \phi_{12} - (x_1 - x_2) \phi_{13}}{k[(x_1 - x_3)(y_1 - y_2) - (x_1 - x_2)(y_1 - y_3)]},$$

$$\cos \alpha = \frac{(y_1 - y_2) \phi_{13} - (y_1 - y_3) \phi_{12}}{k[(x_1 - x_3)(y_1 - y_2) - (x_1 - x_2)(y_1 - y_3)]}.$$
 (A-4)

Since k is always positive, consideration need only be given to the other terms. Letting D stand for the quantity in square brackets, the direction, α , for D \neq 0 is given by:

$$\alpha = \tan^{-1} \left\{ \frac{\left[(x_1 - x_3) \, \phi_{12} - (x_1 - x_2) \, \phi_{13} \right] / D}{\left[(y_1 - y_2) \, \phi_{13} - (y_1 - y_3) \, \phi_{12} \right] / D} \right\}. \tag{A-5}$$

A unique value for direction can be obtained by considering the signs of both numerator and denominator. The quantity D differs from zero for all nonlinear arrays as shown below.

Let $x_1 = y_1 = 0$, D will equal zero for $y_2/x_2 = y_3/x_3$; thus, x_2 , y_2 and x_3 , y_3 will be on a straight line with slope given by this ratio.

APPENDIX B

COMPUTER OUTPUT FOR CROSS-SPECTRA COMPUTATIONS

Table. Guide to computer output.

	1 41	ore. Gurde to computer output.
Line	Column	Explanation
(from top)	(from left)	•
1		Title, plus date and time of the observation (day, month, year, hour, and minutes).
2 to 4		Headings for columns. The numbers separated by dashes in the fourth line give the numbers of the gages in the array (see Fig. 3).
	1	Sequential number of bands 0.0107 hertz wide.
	2	Period at center of band (seconds).
	3 to 7	Percent of energy in each band for gages 1 to 5. Normalized to the energy content from approximately 30 to 3 seconds.
ar ::::::::============	8 to 17	Resulting "representative" direction of wave propagation for the 10 arrays for the corresponding band (degrees).

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ULGREES	1-3-5	-53	137	-17	22	7	-	•	•	1	ď	-1	150	156	152	47.	- 37	-37	- 58	- 50	٠.٠	- 36	07.	65.	1 59	97-	- 76	œ ;	-26	9	- 28	1 39	7 + 3	105	7
CTION (1-5-4	-138	50	10	30	t	17	.5	<u>۷</u>	¥	:	0	7	7	D -	ē	- 30	1	41.	 	10 3 1	1 1 1	.50	1	133	7011	-87	131	- 19	- 60	5,1	4	071	50	1 40
PUSITI	1-2-5	•155	125	7	30	7	1,4	1.5	2	n	0.7	•	170	100	163	161	447	147	150	1 4 5	-170	101	•135	-17	٧,	-	-172	158	9	•	157	-41	133	₹.	-13
	1-5-6	-160	3	-27	77	10	0≥	-	52	7	7.7	0	-		27	9 Z	3.7	30	Ş	0.2	7		=	,	171	-177	151-	<u>ل</u> د		•	113	• . t	1 4 1	1	401-
	1-2-3	-174	34	0.7.	55	13	23	91	2	3	-	•	₹.	er T	3	•	170	173	174	174	-178	170	-176	-176	~	•179	-177	45	170	171	177	7	ů,	<u>.</u>	7
		1.46	• 10	•5•	.3.	.87	18.07	42.54	5.54	25.5	2.24	\$6.	٠,	1.10	. 70	. 50	, ≥ 4	. 43	,6.	1.53		-62	77.	~ 0	.10	5 0.	2 0.	70.	10.	00.	00.	٠٥٠	၁၃•	00.	00.
CTBA		1.25	20.	٠2٠	٠٤٠	•0•	34.34	30.08	D C	3.21	1.29	1.15		.75	.70	90	.39	1.35	·	1.73	.57	.53	.3.	٠١٠	90	₹0.	70.	10.	10.	•	00.	· 0	00.	00.	00.
MESSUME SPECTR		۶,	.07	~.	٤٤.	. 67	70.00	37.03	*0.0	, ¥. ~	. 65	, e 5	95	J.	.71	7	3.5.	1.22	1,20	1.65	, t.	.59	۲,۰	-,	•	90	•	10.	10.	10.	0°°	90.	00.	00.	0.
23.7		٠٢.	~.	. 21	٥,	?	36.44	41.70	00.0	۶۰۰۵	٥٠٠	ζ,	. 26	•	•	. 30	۲۶.	1.2	1.00	,	.20	•\$•	۲۶.	~	=	•0•	5 0.	₹.j.	.01	٠ •	> •	၁ ၀ •	٠ •	10.	0
		1.03	1.10	22.	٠\$٥	, .	00.07	34.15	9.31	2.7.	39.1	1.28		\$6.	0		æ ₹	1.61	1.27	6.37	. 7	۲7.	~ ,	.53		90.	50.	₹0•	.0.	.0.	00.	00.	0	00.	00.
P£#100	したいたる	170.07	*2°04	50.57	20,20	4.°C?	16.79	14.22	12.3.	0	4.75	9.63	90.0	۲.۰۰	5.97	0	66.5	5.03	5.51	20.5	4.1	.,53	. 32	٠.٠	3,95		30.0	15.5	3,36	3.26	3.15	3.05	2.05		2.18
0440	<u> </u>		~	_	3	s	•		⋖,	•	<u>-</u>	=	~	-	3	15	<u>-</u>	_	•	<u> </u>	0	≂:	≎:	5	₹:	€.	9	~	9~	2	S .	7	≃:	٠,	3

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94 ON PE	#100		PHES	PRESSURE SPECTR	CTRA				DIRECTION C	C110% (CUEGHLES	Σ 2 3	SEA A H	NOW AL	7	
.	ENTER						10,01	9-6-1	P05171	VE VALU	ES ARE	COUNTER	(- OC + -	186	;	•
}					•					3 - 5		· •	7-7-7	6.3. 3	V-7-7	5-6-9
- 1	70.67	1.73	1.95		. 93	1.28	-176	٠.	• : •	9		• 32	101	32	9	-100
• ·	9.0	1.28	000	•	1.07	9:1	9	Š	07	75	0,7	30		10	2	9
~ ·	6.57	1.0k	21.1	.01	•	1.01	-171	. 161	-100	-139	9	101	150	173	-174	.167
	9:50	2		٥.	\$ 7.	٥2.	136	102	117	0	5	00	105	120	308	100
· ·	91	20	9.	0	.67	•	37	<u>-</u>	=	40	•	53	•	2	•	52
•	6.79	30.00	×1.45	26.39	20.05	24.64	2	2.5	7	25	13	30	• •	26	2	2
- 1	4.22	15.58	000	15.58	12.75	15.62	23	=	82	2	₹	35	2	28	2	
·	2.34	70.4	9.70	4.1.	14.60	10.51	7.	2	7	~	•	7	. %	\$2	9	2
'	0.00	4.33	3,15	3.75	3.66	2.67	•	~	3	0	î	-	~	2	•	2
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<u></u>	7.42	. 6	ē.	100	\$6°	. 19	-13	-	-	-	-15	~	• 154	-	9	• ;
3 (6.87	9	7.60	1.71	.67	1.00	-	20	145	- 41	- 30	-37	-151	- 163	. 101	•
<u>~</u> :	010	2.4	00.5	₽ O.	50.5	3.26	172	35	155	.50	~ 7 .	, y	,	135	-152	97
4.		20.6	9.7	50.5	9	6.50	174	7	156	*5		30	97-	•133	-150	37.
		77.5	9.00	6.27	5.61	9.51	174	2	150	-50	-37	07-	-1.5	-131	-147	503
0 6	17:51	7.00	10.66	7.7	7.12	÷	173	33	143	3 1	- 15	• 35		1.20	79.	
	20°5			4.05	4.37	4.39	1,76	-	100	931		- 43	0 7 7	-125	-135	100
	•	0.5	7.37	9,	70.7	2.70	170	0	179	-54	9	• 50	7	103	-150	001
		1.56	3		•	1.61	S.		-	105	166	70.	=	63	-141	-173
			3 4		٠. د د	1.00	-173	-23	7.	• 56	• 50	9.	101	=	117	137
					.37		٠ <u>٠</u>	?	-15	3	150	•\$	3. -	4	7.7	103
	1.70		. 1	ĵ.	3 c	6	<u>-</u>	?	. 6	4	20	-53	78	~	3	70
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	10.1			9 4	•	-	-1,1	-15	~ ?	26.	. 53	12.	5	2	55	-139
		•				ζ.	0	176	2.	163	126	-121	¥ (? •	98-	°	-103
	2.5	•		•		• 0	173	777	5 = 2	100	- 36	1 45	-47	3	51	•
	9346			•	.0.	~0.	105	-	Ş	-21	02-	-		-155	-131	00
		•			.01	-0	0	-170	-106	120	1 5	151	-151	-110	~>	0 > -
; 2	2.05	•		•		- -	177	2	20	3	5.	3.	-163	-121	-1.0	7.5
	F 5 4 3	•	2	3 6	10.	30.	~	-27	* ? C *	÷	~ eb	7 a .	55	74	~	047-
1 3	2.78	- 0) :) :	00.	00.	~	27.	-	5	0 7	٥	3.5	0.0	52	3
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		3-4-5	171	; ;	•	~~	27	2	52	25	52	52	=	5	-20	• 1 •	-28	-2	103	70	102	102	6	~	-145	1771	25	3	- 6	- 4	-15	-85	-87	7771	125	~
	_	5-4-2	-157	2,	7	~!-	*2	82	54	63	₹,	9	۲>	1.0	,	3	-153	-150	-147	-150	-153	-166	-121	•115	7.5	3	Š	74	27.	?	•	•	129	108	57	3
	NOWMAL	2-3-5	-171	3	20	-	<u>-</u>	\$	5 2	25	~	92	€	15	-21	97-	-136	-133	.131	-126	100	ô	9	 Đ	7.	54	7	8≥	~	- 58	9	-131	6	115	52	58
	SEAMAND	2-3-4	172	~ ,	3.	?	0	۶,	۲,	25	≈	97	٠ ج	15	101	6711	-114	0,7.	-145	₹ 7	3.5	30	٥,	2	0	24	25	13	٤,	-123	07-	-155	7	5.3	53	33
	CONTER	1-4-5	179	771	=	10.	9	35	2	2	5	36	34	-54	- 35	- 34	• 34	-37	-37	27-	0 7 1	3 -	-54	-21	136	95-	176	101	0 7	143	-100	120	127	707	123	191
	DEGREES	1-1-5	-173	- 37	\$ \$	- 42		7	2	•	•	0	î	155	- 35	15-	- 58	- 57	- 57	- 50	- 59	-	-50	- 25	- 75	0 7	-172	- 34	7.7	143	57.	- 52	- 51	-62	143	e, .
	CT TON C	3-3-4	-170	-171	23	-113	%	52	<u>-</u>	1.7	2.	7	3	10	3 1	-35	1 6	7 7 1	167	117	1 2 1	1 6	• 5e	-50	•135	۰,	-170	-110	22	131	-101	190	- 93	-150	123	131
	POSITI	1-2-5	172	971	9	•53	•	7.7	<u>-</u>	•	2	72	₹	170	- 55	142	1 4 7	4	7 = 0	155	155	170	-156	• 13	02-	-83	159	02.	152	1 4 3	155	103	25	-173	î	-71
		1-5-1	171	165	07	-124	25	27	2	~	~	27	52	-	- 38	56	3,6	t Q	9	٥,	1.7	7	• 2 1	-	-151	•33	1 6 9	-151	9₹	156	136	137	152	-176	-152	-135
		1-5-3	174	172	74	-147	7	<u></u>	7.2	52	\$	9	9	۲>	57-	101	171	171	172	174	175	178	-175	-177	001.	•	101	-174	95	75	172	174	170	-177	•	•
			6.	•••	.38	.13	97.	8.72	40.04	10.58	3.47	.74	.57	00.	5.03	3.10	B.12	7.15	2.53		3.0 0	2.0b	1.24	• > 6	•	٠,	.23	>	50.	50.	.01	10.	00.	00.	30.	00.
	CTRA		1.15	.5	.34	.07	.37	10.34	****	90.0	5.51	.60	.74		1.71	3.07	5.99	4.17	3.71	3.59	2.34		٠,٢٠	• : •	.17	• • •	90.	•	.01	.03	10.	.01	.01	00.	00.	•
2108	PHESSURE SPECTR		. 97	\$9.	.35	20.		9	37.50	10.70	3.41	7	. 75	.63	5. 00	3.50	26.9	9.63	00.7	3.47	2.31		\$ 9.	1,32	٠.	. 24	40.	60.	90.	₹0.	.01	7 0.	00.	3.	9 0.	2
5 6 1970	PRES		1.04	.87	3.	.15	• 20	6°°6	35.24	15.45	3.71	7	.73	•	30.	3.44	7.86	12.8	3.72	0 0	5.65	- 60	1.35		= :	~		•	50.	3	90.	.01	.01	•	•	99.
SE ARHAY 2			.77	9 3 .	6≥•	.01	۲ŋ.	10.57	50.72	4.27	2.53	٠,	- -	.50	97.	3.75	2.01	6.57	9	٥٠٠	5.03		7:-	00.	٠.	9.	<u>.</u>	\$0.	~0.	?	~ 0•		10.	00.	•	00.
PT MUGU FIVE GAGE	PE-130	E>169	170.07	*6.34	30.57	65.65	S: 6	19.79	125	12.3.	37.30	6,75	6,33	400	75	0.17	0,.0	9.0	5.93		۶۰5 ۶	٠,٠	53	32		\$ · •	, , o	3,5	3.51	9.10	3.26	3.15	3.75	5 2	2.4	3. 7€
PT 706	BAND	J	-	~	-	3	ď	•	~	•	•	2	=	~!	~	-	15	-	-	•	-	20	7	2	: S	2	€;	9	۲,	97	2	9	-	2	ς,	3

		3-4-5	174	142	-	.0-	~	52	22	•	22	80	3	?	-38	-23	- 32	-31	100	102	80	101	717	•	7	771-	154	-155	• : : •	19-	2	-	=	=	-23	=
		2-4-5	20	- 10	51.	72.	0 7	Ş	22	•	53	77	71.	-15	9	-100	-154	-152	911	-1+2	-135	-132	-147	-121	g 2	0	110	Ð	75	-130	-130	170	100	116	5,	9
	NORMAL)	-3-5 2	20	101	-	-18	01	52	2	9.	23	20	?	~	. 32	-140	-130	-134	• : 32	-156	-123	101	2.0	4	67	64	117	24	•	-158	-140	-143	-129	9	.73	-5-
	5544400 CLOCA418	2-3-6	101	137	1.3	-15	0.7	52	23	1.0	5 2	97	~	•	-150	-150	-1.7	1111	-1-5	-1.3		25	62	€ 2	305	۲,	1.55	9,	\$?	0.1	-19	17.0		;	-1.6	;
	COUNTER	1-6-5	129	170	26	-116	75	3	£	10	30	21	•	-20	770	*2·	• 39	• 30	• 30	*38	-36	.53	-47	3 7	17	155	151	101	175	- 34	127	3	~	65	-61	9
	DEGREES ES ANE	1-5-5	97-	-167	-17	179	P	3	-	0	0	•	97.	77.	771	15.	- 38	- 36	-37	-37	- 34	-65	- 4	2 4 5	135	-15	.0.	61.	9	145	-29	67-	•	~	47	1,58
	CTION C	1-3-4	-136	-179	-15	-151	2	•	≈	1	=	3.		-15	- 43	-35	1 45	- 6 2	-45	143	77-	071	3	1 40	93	155	-168	135	-121	17	-104	~ * *	~	-	100	127
	0146	1-2-5	- 1 4	175	•	1011	3	15	2	=	-	ຈ	-17	- 45	531	-30	1.4	777	148	971	134	179	-107	-101	-20	9	117	174	00.	-157	107	- 23	-21	105	162	150
		1-5-1	-155	175	•	-152	•	2	7.7	1.5	•	80	-20	-25		17.	5.8	51	77	5.5	23	~	-1	î	- 36	138	178	175	-130	7-	137	-37	.3	7	•	153
		1-5-1	-167	176	~	-152	^	;	92	1.7	%	62	-	• ? •	071	- 57	170	170	172	173	174	179	-177	-177	£.	150	179	-	-101	7	171	-70	-12	154	~	150
			96.	55.	•50	.56	.76	3.23	15.00	3.46	4.25	4.45	1.21	76.	1.21	1.10	40.71	10.62	12.07	3.08	2.32	3,37	63.	٧,٠	.35	٥٢.	• 28	=	10.	₹0.	.01	00.	 -	10.	3	0°0°
	CTRA		1.65	÷.	.36	.54	٠,	2.87	13.58	3.40	90.9	1.55	. oe	1.36	1.54	7.61	24.63	14.75	6.67	4.34	• : -	1.17	.36	6	.75	•	.30		.07	.03	20 •	.0.	· •	00.	00.	0
1970 1758	KESSURE SPECTRA		1,93	7.	. 47		.01	3,45	13.00	3.40	6.25	2.34	1.73	1.10	1.51	7.30	58.19	14.67	6,29	4.30	00.1	" "	0	•	. 7.7	?:	; :	•	50.	٠٥٠	50.	•	-	•	ទ	•
0 197	PKES		2.10	٠.	9.	٤٤.	2.00	70.0	12.20		5.27	1.05	. 0.	1.01	•	7.63	28.82	11.47	9.78	3,24	27.	2,36	¢.	•	•	-	0	. 23	•	.0			6	70	3	
E ARHAY 2			4.	.51	٠٠,	7.		3.70	19.11	P.04	0.0	•	1.20	ę.	3.65	7.34	<3.e3	10.00	£ . 70	5.25	1.32	1.70	2 v .	1.20	9 0	•	•	02.	• 0 •	70	• 0 •	90.	•	•	00.	•
PT MUGU FIVE GAGE ARMI	ER 100	CLWTED	170.67	60.24	36.57	20.26	20.48	16.79	14.22	12.34	10.84	5.	8.83	90.0	7 · " >	6.87	9.0	5.40	5,63	5.31	20.5	6.70	 	27.		٠,	0,	30.0	1,,,	96.5	97.5	51.5	50.4	5,42	9617	9, 1
PT MUG	BAND PERIOD	ŭ	-	~		3		•		 •0 (2		3 (5	•	~ .	Ð (-	20	7	> :	S		Ç :	? ;	2	D	.	3		× :	3 -	;

PT MUGU FIVE GAGE ARMAY 28 & 1970 2010 BAND PERIOD PRESSUR SPECTA

027	PER100		PHES	PRESSURE SPECTE	CTRA				DIPE	1 10 C	DEGREES ARE	F POR	SEAMABO	NONTAL	_	
	CENTER						1-2-3	1-5-1	1-2-5	1-3-4	5-6-1	1-6-5	7-7-6	2-3-5	5-4-2	3-4-5
-	170.67			•	•	.0.	•	-142	• 36	-159	-36	77	-143	07-	35	•
~	40.24	47.	75.	97.		97.	. 65	=======================================	101-	-125	-1.4	-	-::-	76.		-121
~	35.57	.78	1.12	\$9.	0.	•	•110	4	7/-	-11	-58	- 1	• 55	- 70	001	-
•	26.26	.7.	.61		ē.	.00	-	•	-75	- 69		-95	- 92	99.	• > •	-72
•	80.08	1.40	9.	9.	.67	1.03	103	9	9	52	~	25	0,7	9	77	7
•	16.74	. 69	1.57	1.75	1.50	1.38	•	2	71	•	•	€	<u>.</u>	:	2	2
~	14.22	3.65	12.0	4.70	2.07	4.67	~	12	2	12	3	7.	•	0.7	•	9.
•	12.34	Z.00	1.01	₹.03	2.51	2.05	53	77	::	'n	3	36	7	7	7	7.
•	10.69		1.20	1.10	88.	1.11	0	•	•		-1	\$2	7	•	7	7
2	9.75	1.10	Ŷ.	1.03	1.09	.71	• .	-15	-15	•	• 7 •	~	~	•	1	0
=	6.83	7.60	7.80	6.67	7.7	. O	-37	٥,	-24	-32	-27	• 20	9.7	•15	71.	7.
~	90.8	16.02	19.60	19.80	20.27	60.27	07	• 33	62-	• 30	• 30	-~	-152	~~	7	-2-
3	7.42	10.78	9.17	0.50	9.54	4.67	67-		• 5 •	- 35	- 30	-27	-152	-20	9	-20
=	6.87	14.31	17.51	10.45	14.77	17.77	071	- 34	67-	•5-	-30	-27	-151	~~	4	-21
=	0.40	12.96	10.10	10.39	11.64	9.70	6 3 •	• 38	• 33	10	- 34	•33	971-	-138	-157	•25
=	2.00	9.71	4.67	7.39	7.54	*1.0	00.	25	130	- 30	. 31	٠.	.148	-136	-155	-23
_	5.63	1.95	9.5	4.39	4.33	6.70	1 70	43	120	-37	- 50	-25	-145	-130	071.	701
=	5.31	4.37	1.0	2.61	5.24	2.06	172	2	128	0 7 1	- 52	92.	241.	• : 26	-140	9.
•	20.5	2.53	9.4	3.57	3.10	e	175	2	159	071	9	. 4	~ 1	-154	-:38	•
2	. 7	2.9	 	2.03	1.47	1.00	176	7	°.	9 1	-29	•	~	•	-125	•172
₹	4.53	•		1.58	1.59	1.51	175	:	142	.38	• 35	- 35	32	63	-145	70
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APPENDIX C

FOURIER COEFFICIENTS FOR A MIXTURE OF THREE SINUSOIDS

The Fourier Transform of the function:

$$f(t) = A \cos(\sigma t - \phi) \tag{C-1}$$

with

$$\sigma = \frac{2\pi(\hat{\mathbf{m}} + \delta)}{N\Lambda t}, \quad |\delta| \le \frac{1}{2}, \qquad (C-2)$$

where 1 << \hat{m} << N, computed from N values of f(t) evaluated at equal increments of t; Δt is given by the set of coefficients (Harris, 1974):

$$a_{m} = \frac{2A \sin \pi \delta \cos(\pi \delta - \phi)}{N} \left[\frac{1}{\tan[\pi(\hat{m} - m + \delta)/N]} + \frac{1}{\tan[\pi(\hat{m} + m + \delta)/N]} \right],$$

$$b_{m} = \frac{2A \sin \pi \delta \sin(\pi \delta - \phi)}{N} \left[\frac{1}{\tan[\pi(\hat{m} - m + \delta)/N]} - \frac{1}{\tan[\pi(\hat{m} + m + \delta)/N]} \right],$$

$$m = 1, 2, \dots, \frac{N}{2}.$$
(C-3)

Harris shows that for values of m near \hat{m} , and for \hat{m} far removed from 1 and N/2,

$$a_{m} = \frac{A \sin \pi \delta \cos (\phi - \pi \delta)}{\pi (\hat{m} - m + \delta)},$$

$$b_{m} = \frac{A \sin \pi \delta \sin (\phi - \pi \delta)}{\pi (\hat{m} - m + \delta)}, \quad m = 1, 2, \dots, \frac{N}{2} \quad (C-4)$$

are good approximations to the coefficients in equation (C-3). Equation (C-4) shows that convergence is slow.

For the special case δ = 0, substitution into equations (C-3) or (C-4) gives:

$$a_m = b_m = 0, \qquad m \neq \hat{m} ,$$

and

 $a_m = b_m = indeterminate for m = \hat{m}$.

Use of L' Hospital Rule in equation (C-3) when $\delta \rightarrow 0$, shows that:

$$a_m = A \cos \phi$$
, $b_m = A \sin \phi$, $m = \hat{m}$. (C-5)

Application of the cosine bell data window,

$$\mathbf{\hat{f}}(n\Delta t) = \frac{1}{2} \left\{ \left[1 - \cos \frac{2\pi n}{N} \right] \right\} \mathbf{f}(n\Delta t) , \quad n = 1, \ldots, N , \quad (C-6)$$

is equivalent to replacing the original sinusoid $f(n\Delta t)$ with the sum of three sinusoids (Harris, 1974), where

$$\stackrel{\sim}{\mathbf{f}}(\mathbf{n}\Delta\mathbf{t}) = \frac{A}{2} \left[2 \cos \left[\frac{2\pi \left(\hat{\mathbf{m}} + \delta \right) \mathbf{n}}{N} - \phi \right] \right] - \cos \left[\frac{2\pi \left(\hat{\mathbf{m}} - 1 + \delta \right) \mathbf{n}}{N} - \phi \right] - \cos \left[\frac{2\pi \left(\hat{\mathbf{m}} + 1 + \delta \right)}{N} - \phi \right] \right] .$$
(C-7)

For δ = 0, and in view of equation (C-5), the Fourier Transform of this modified function will be given by:

$$a_{\hat{m}-1} = a_{\hat{m}+1} = -\frac{A}{4}\cos \phi$$
, $b_{\hat{m}-1} = b_{\hat{m}+1} = -\frac{A}{4}\sin \phi$

and

$$a\hat{m} = \frac{A}{2}\cos \phi$$
, $b\hat{m} = \frac{A}{2}\sin \phi$

with $a_m = b_m = 0$ for all other values of m. Thus, energy appears at three adjacent m values.

Harris (1974) shows that after application of the cosine bell data window, the approximate values of the coefficients are given by:

$$a_{m} \doteq \frac{A \sin \pi \delta \cos(\phi - \pi \delta)}{2\pi (\hat{m} - m + \delta) [(\hat{m} - m + \delta)^{2} - 1]},$$

$$b_{m} \doteq \frac{A \sin \pi \delta \sin(\phi - \pi \delta)}{2\pi (\hat{m} - m + \delta) [(\hat{m} - m + \delta)^{2} - 1]}, \qquad m = 1, \dots, \frac{N}{2}. (C-8)$$

Thus, convergence increases rapidly, and values of the coefficients for $(\hat{m} - m) \ge 3$ may be disregarded.

Equations (C-3), (C-4), and (C-8) imply that

$$\tan\left(\phi - \pi\delta\right) = \frac{a_{nl}}{b_{ml}} \tag{C-9}$$

Since δ may be as large as 1/2, phase values computed from the Finite Fourier Transform may be in error by as much as 90° .

For a simplistic simulation of a wave train, it is sufficient to combine three sinusoids with nearby periods propagating in the same direction. Letting A_i equal the amplitudes and k_i , i-1,2,3, the wave numbers, the Fourier Transform of the combination is given by:

$$a_{m} = \sum_{i=1}^{3} \frac{A_{i} \sin \pi \delta_{i} \cos (\Phi_{i} - \pi \delta_{i})}{2\pi (\hat{m}_{i} - m + \delta_{i}) [(\hat{m}_{i} - m + \delta_{i})^{2} - 1]},$$

$$b_{m} = \sum_{i=1}^{3} \frac{A_{i} \sin \pi \delta_{i} \sin (\Phi_{i} - \pi \delta_{i})}{2\pi (\hat{m}_{i} - m + \delta_{i}) [(\hat{m}_{i} - m + \delta_{i})^{2} - 1]}.$$
(C-10)

Nearby periods are attained by setting

$$\Delta_{i} = (\hat{m}_{i} - m) \leq 3. \qquad (C-11)$$

Let the coordinates of three nearby locations be (x_j,y_j) ; j=1,2,3. The only difference among the Fourier Transforms (eq. C-10) arising from wave records at each location is the values of the Φ_i 's. At each location j, the Φ_i values are:

$$\Phi_{ij} = k_i(x_j \cos \alpha + y_j \sin \alpha) - \phi_i$$
, $i = 1,2,3$, (C-12)

or

$$\Phi ij = k_i \Omega j - \phi i ,$$

where

$$\Omega_{j} = x_{j} \cos \alpha + y_{j} \sin \alpha$$
, $j = 1,2,3$. (C-13)

Since the three sinusoids are assumed to have nearby periods, let

$$k_1 = k_2 = k_3 = k$$
.

Inus,

$$\Phi_{ij} = k\Omega_j - \phi_i.$$

Then, for the wave record at location j:

$$a_{mj} \doteq \sum_{i=1}^{3} \frac{A_{i} \sin \pi \delta_{i} \cos (k\Omega_{j} - \pi \delta_{i} - \phi_{i})}{2\pi (\Delta_{i} + \delta_{i}) [(\Delta_{i} + \delta_{i})^{2} - 1]},$$

$$b_{mj} \doteq \sum_{i=1}^{3} \frac{A_{i} \sin \pi \delta_{i} \sin (k\Omega_{j} - \pi \delta_{i} - \phi_{i})}{2\pi (\Delta_{i} + \delta_{i}) [(\Delta_{i} + \delta_{i})^{2} - 1]}.$$
(C-14)

Let:

[i] =
$$2\pi(\Delta_f + \Delta_f)[(\Delta_f + \Delta_f)^2 - 1]$$
, i = 1,2,3. (C-15)

The expanded expression for $a_{m,i}$, after collecting terms in $\cos k$; and $\sin k \beta_i$, is:

$$a_{rij} = \cos k\pi_{ij} \left\{ \frac{A_{1} \sin \pi\delta_{1} \cos(\phi_{1} + \pi\delta_{1})}{[1]} + \frac{A_{2} \sin \pi\delta_{2} \cos(\phi_{2} + \pi\delta_{2})}{[2]} + \frac{A_{3} \sin \pi\delta_{3} \cos(\phi_{3} + \pi\delta_{3})}{[3]} \right\} + \sin k\psi_{ij} \left\{ \frac{A_{1} \sin \pi\delta_{1} \sin(\phi_{1} + \pi\delta_{1})}{[1]} + \frac{A_{2} \sin \pi\delta_{2} \sin(\phi_{2} + \pi\delta_{2})}{[2]} + \frac{A_{3} \sin \pi\delta_{3} \sin(\phi_{3} + \pi\delta_{3})}{[3]} \right\};$$

and similarly for b_{p_2} :

$$b_{mi} = \sin k a_{ij} \left\{ \frac{A_{1} \sin \pi a_{1} \cos (\phi_{1} + \pi \delta_{1})}{[1]} + \frac{A_{2} \sin \pi a_{2} \cos (\phi_{2} + \pi a_{2})}{[2]} + \frac{A_{3} \sin \pi a_{3} \cos (\phi_{3} + \pi a_{3})}{[3]} \right\}$$

$$- \cos k a_{ij} \left\{ \frac{A_{1} \sin \pi a_{1} \sin (\phi_{1} + \pi a_{1})}{[1]} + \frac{A_{2} \sin \pi a_{1} \sin (\phi_{2} + \pi a_{2})}{[2]} + \frac{A_{3} \sin \pi a_{3} \sin (\phi_{3} + \pi a_{3})}{[3]} \right\}.$$

For every narrow-banded wave train, $|A_{\uparrow}| \ge 3$ in equation (C-15), unpredictable terms are introduced in the expressions for the coefficients

en jar 1911 jarosiaj which make the ratio b_m/a_m a poor estimator of the value of the phase of the sinusoid.

The numerators of the terms inside the braces in the above equations for a_{m_i} and b_{m_i} are at most of order A_i , i=1,2,3.

Randomness in the values of $\circ_{\mathcal{S}}$, i=1,2,3 and in the phase relationships of the three sinusoids at the origin of coordinates might produce partial cancellations among the terms inside the braces to reduce the resulting error.

Assume for example: $\phi_{i} = 0$, i = 1,2,3. The coefficients reduce to:

$$a_{mj} \stackrel{:}{=} \frac{1}{2} \cos k\Omega_{j} \left\{ \frac{A_{1} \sin 2\pi A_{1}}{[1]} + \frac{A_{2} \sin 2\pi A_{2}}{[2]} + \frac{A_{3} \sin 2\pi A_{3}}{[5]} \right\}$$

$$+ \sin k\Omega_{j} \left\{ \frac{A_{1} \sin^{2} \pi A_{1}}{[1]} + \frac{A_{2} \sin^{2} \pi A_{2}}{[2]} + \frac{A_{3} \sin^{2} \pi A_{3}}{[5]} \right\} ;$$

similarly, for box:

$$b_{pq} = \frac{1}{2} \sin k(q) \left\{ \frac{A_1 \sin 2\pi\delta_1}{[1]} + \frac{A_2 \sin 2\pi\delta_2}{[2]} + \frac{A_3 \sin 2\pi\delta_3}{[3]} \right\} - \cos k(q) \left\{ \frac{A_1 \sin^2 \pi\delta_1}{[1]} + \frac{A_2 \sin^2 \pi\delta_2}{[2]} + \frac{A_3 \sin^2 \pi\delta_3}{[3]} \right\}.$$
 (C-16)

Letting:

then:

$$a_{nj} = \frac{1}{2} \cos k k_{ij} \left[\frac{L_{1}}{[1]} + \frac{M_{1}}{[2]} + \frac{N_{1}}{[3]} \right] + \sin k_{ij} \left[\frac{L_{2}}{[1]} + \frac{M_{2}}{[2]} + \frac{N_{2}}{[3]} \right] ,$$

and

$$b_{PT_{d}} = \frac{1}{2} \sin k k_{2} \left[\frac{L_{1}}{[1]} + \frac{M_{1}}{[2]} + \frac{N_{1}}{[3]} \right] - \cos k_{2} \left[\frac{L_{2}}{[1]} + \frac{M_{2}}{[2]} + \frac{N_{2}}{[3]} \right]. \quad (C-18)$$

Assume further: $A_1 \sim A_2 \sim A_3$ and $|\hat{m}_i| - m| = 0$ for i = 2 and equal for i = 1 and 3. Since |S| < 1, the terms [1] and [3] are of

approximately the same magnitude but of opposite signs. Thus, terms involving the products [1] [2] and [2] [3] tend to cancel. Letting $\delta_1 = \delta_2 = \delta_3 = \delta$:

$$\frac{b_{mj}}{a_{mj}} \doteq \frac{\sin k\Omega_j \sin 2\pi\delta - 2\cos k\Omega_j \sin^2 \pi\delta}{\cos k\Omega_j \sin 2\pi\delta + 2\sin k\Omega_j \sin^2 \pi\delta}.$$

Using the trigonometric identities for the double arc, this expression reduces to:

$$\frac{b_{mj}}{a_{mj}} \doteq \tan (k\Omega_{j} - \pi c) . \qquad (C-19)$$

Phase differences between locations will be approximately correct.

APPENDIX D

SPECTRA PLOTS AND COMPUTER OUTPUT FOR SIMULATED OBSERVATIONS

Figures D-1 to D-8 show high-resolution spectra pressure gages 1 to 5 at Pt. Mugu, California.

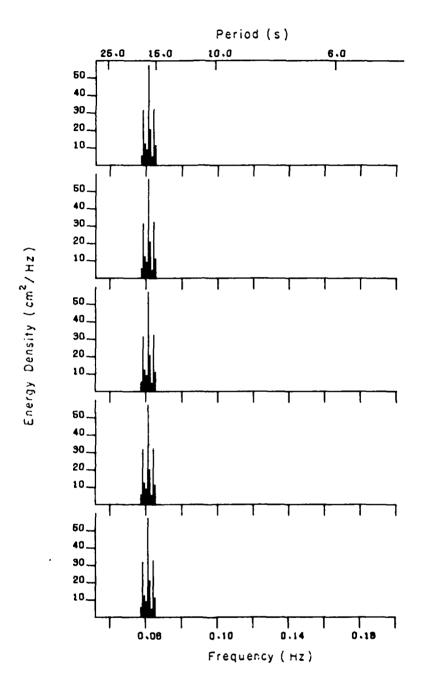


Figure D-1.

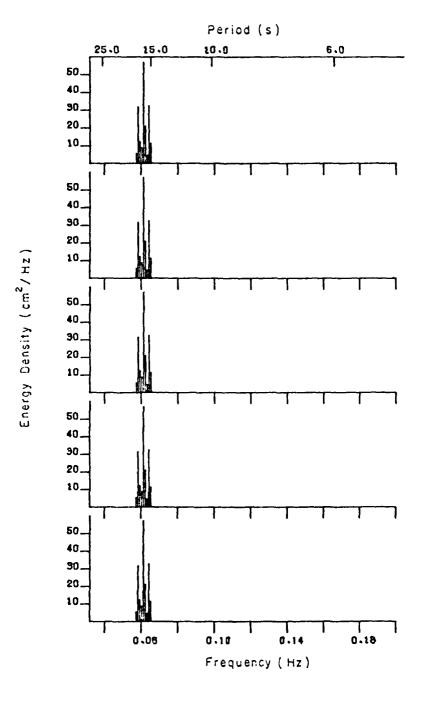


Figure D-2.

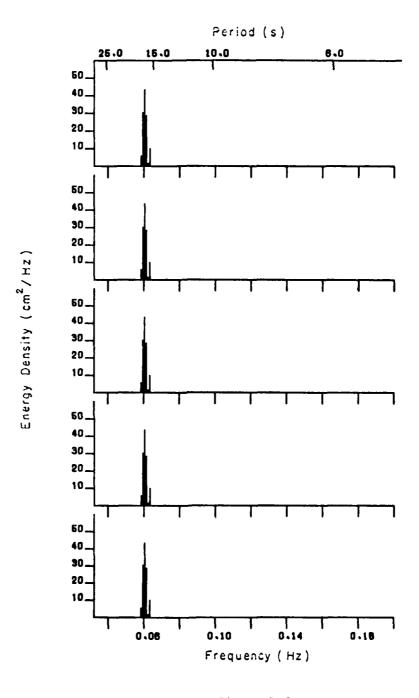


Figure D-3.

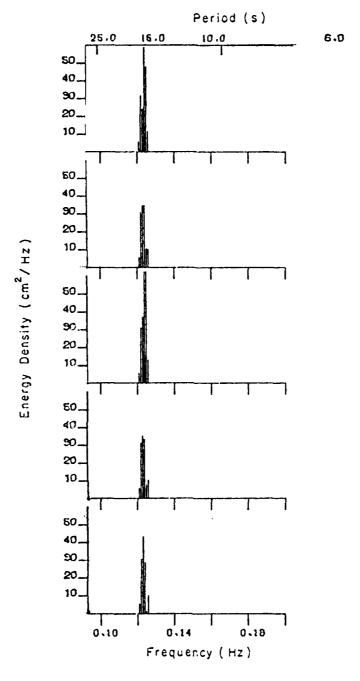


Figure D-4.

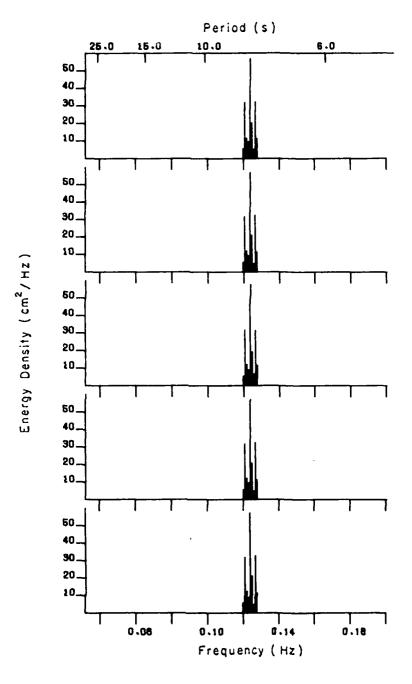


Figure D-5.

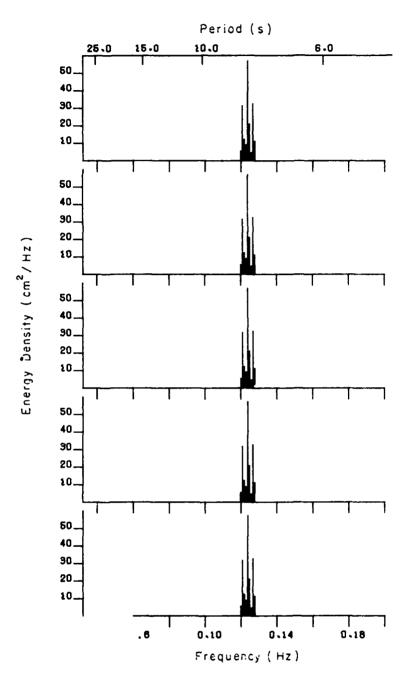


Figure D-6.

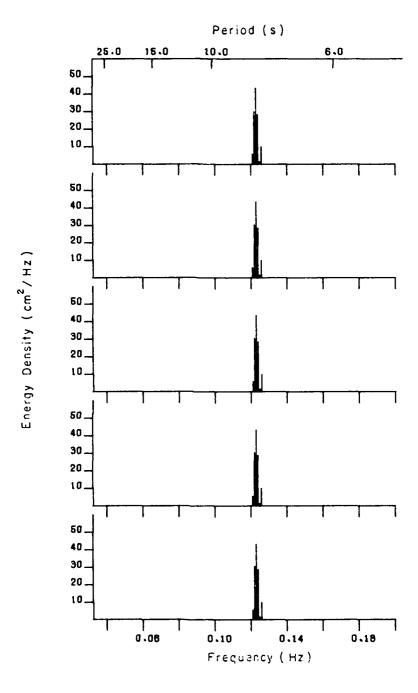


Figure D-7.

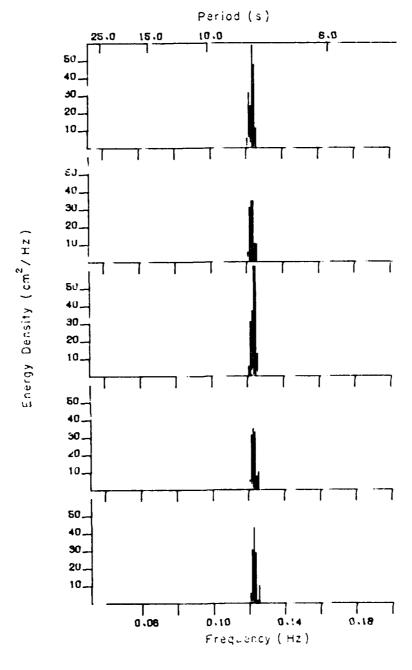


Figure D-8.

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6	PRESGURE SPECTOR (SQ CH)	00.	00.	00.	00.	00.	•	00.	.0.	5.63	30.65	42.15	56.45	43.84	11.27	.02	00.	00.	00.	00.	00.	00.	00.	00.	00.	00.	00.	0	00.	00.	•	0
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Ş	18,29	00.	00.	00.	00.	00.	157	• 153	91.6	• 150	0	-170	•155	.151	-166	-170
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APPENDIX E

HIGH-RESOLUTION SPECTRA FOR FIELD WAVE DATA

Figures E-1 to E-44 show high-resolution spectra for pressure gages 1 to 5 at Pt. Mugu, California. Date and significant wave height are indicated for each figure.

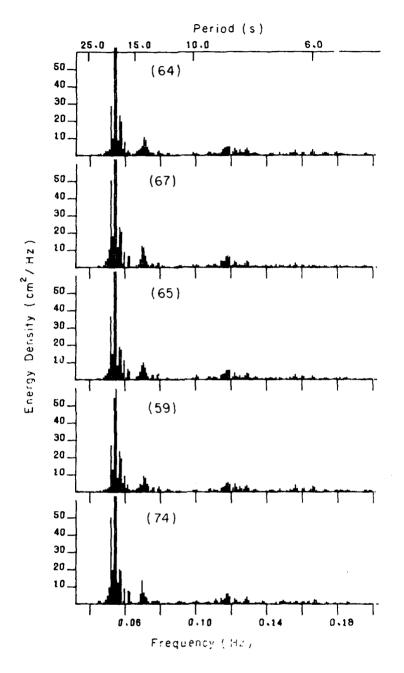


Figure E-1. Significant wave height 93.6 centimeters (3.1 feet), 10 April 1970, 2206 hours.

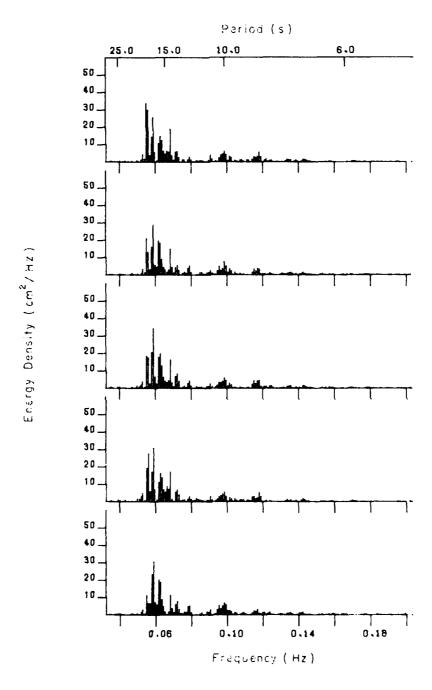


Figure E-2. Significant wave height 84.3 centimeters (2.8 feet), 20 April 1970, 0928 hours

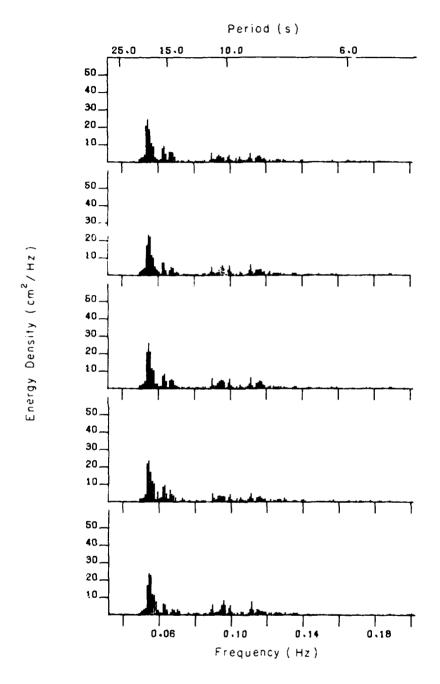


Figure E-3. Significant wave height 81.0 centimeters (2.7 feet), 20 April 1970, 1228 hours.

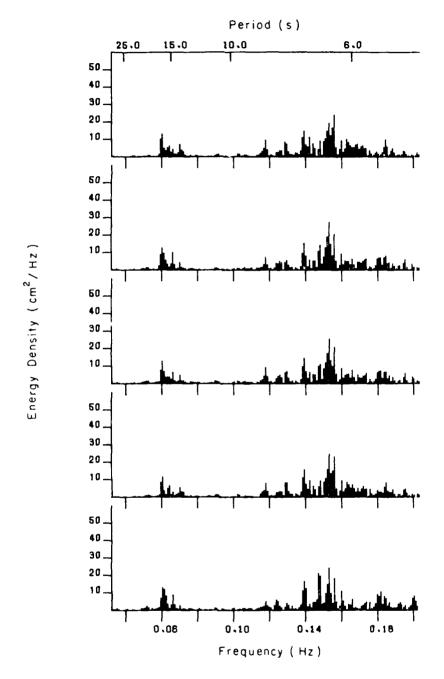


Figure E-4. Significant wave height 105.3 centimeters (3.5 feet), $21~\mathrm{April}~1970$, 0028 hours.

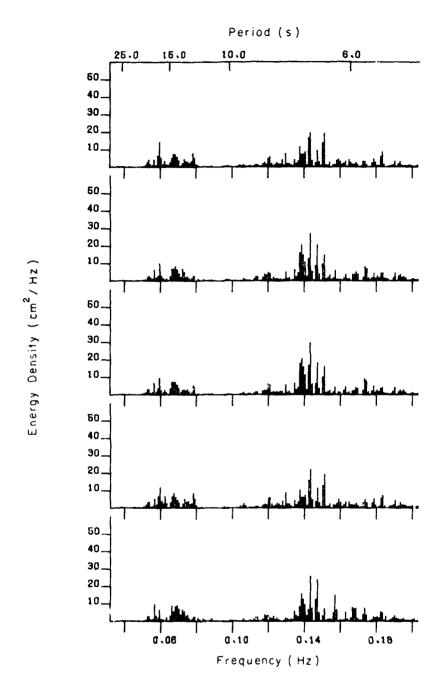


Figure E-5. Significant wave height 95.2 centimeters (3.1 feet), 21 April 1970, 0229 hours.

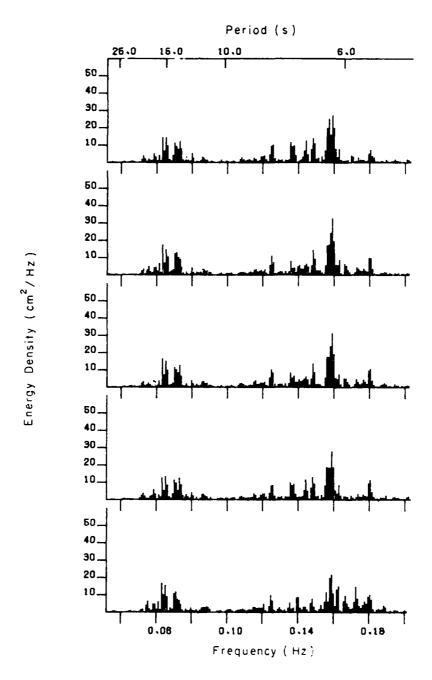


Figure E-6. Significant wave height 100.7 centimeters (3.4 feet), 21 April 1970, 0929 hours.

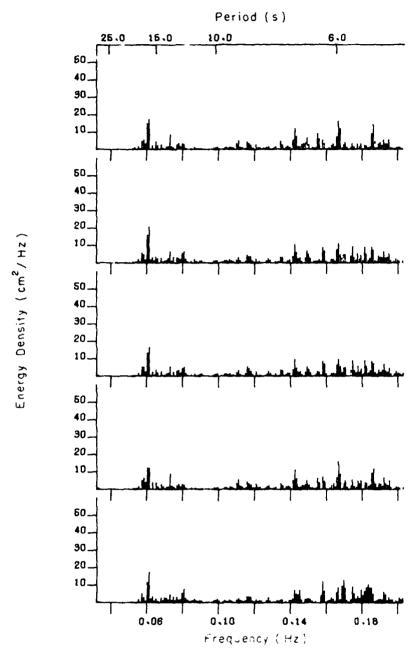


Figure E-7. Significant wave height 94.7 centimeters (3.1 feet), 21 April 1970, 1229 hours.

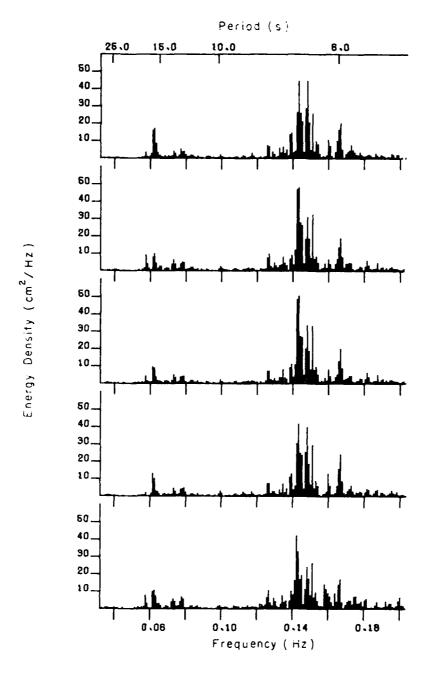


Figure E-8. Significant wave height 113.8 centimeters (3.7 feet), 21 April 1970, 1829 hours.

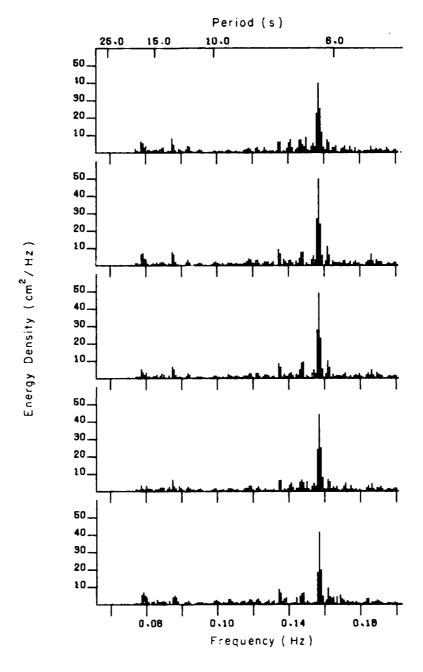


Figure E-9. Significant wave height 85.1 centimeters (2.8 feet), 21 April 1970, 2129 hours.

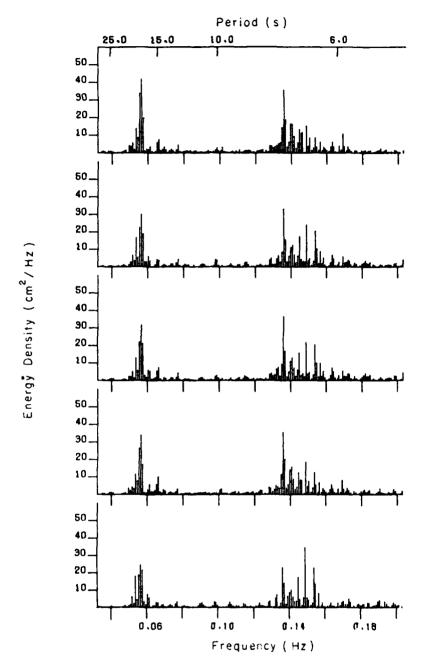


Figure E-10. Significant wave height 109.3 centimeters (3.6 feet), $10~\mathrm{June}~1970$, 1421 hours.

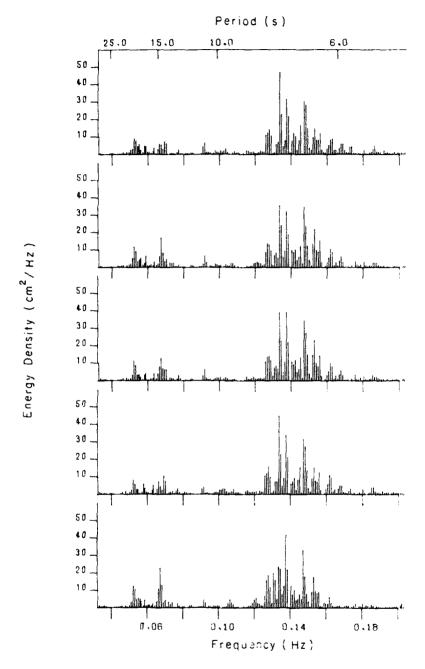


Figure E-11. Significant wave height 117.9 centimeters (5.9 feet), 10 June 1970, 1731 hours.

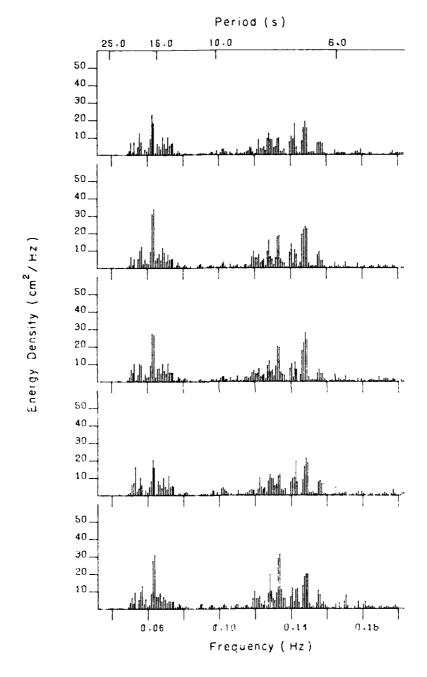


Figure E-12. Significant wave height 113.1 centimeters (3.7 feet), 10 June 1970, 2041 hours.

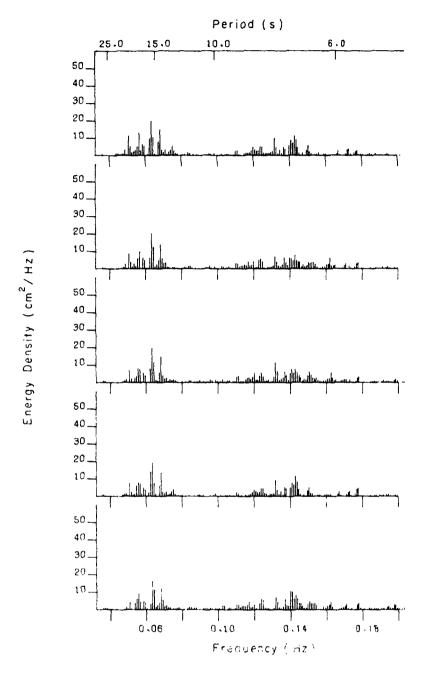


Figure E-13. Significant wave height 87.5 centimeters (2.9 feet), 10 June 1970, 2351 hours.

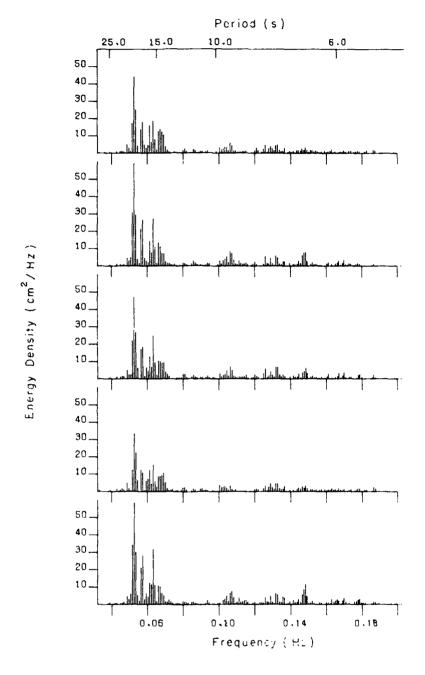


Figure E-14. Significant wave height 93.3 centimeters (3.1 feet), $11~\mathrm{June}~1970$, 0301 hours.

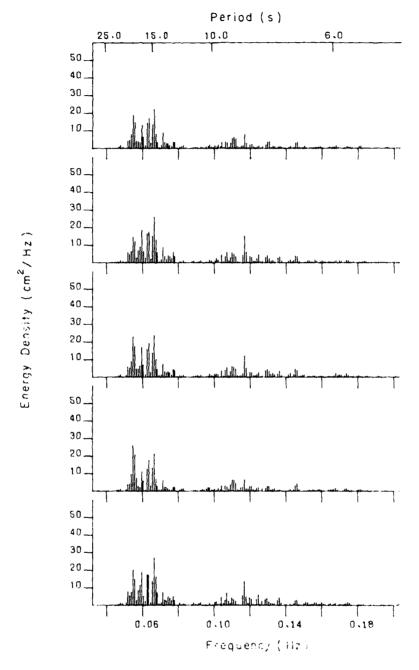


Figure E-15. Significant wave height 81.0 centimeters (2.7 feet), 11 June 1970, 0631 hours.

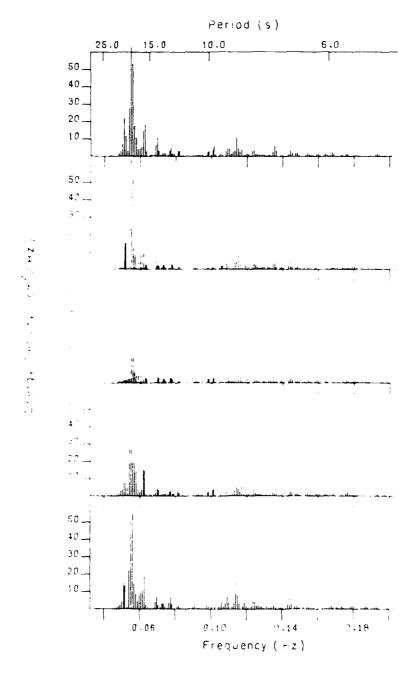


Figure E-16. Significant wave height 88.6 centimeters (2.9 feet). 11 June 1970, 0942 hours.

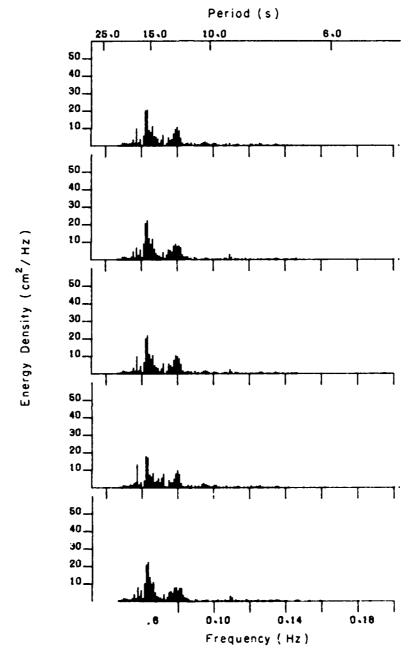


Figure E-17. Significant wave height 71.0 centimeters (2.3 feet), 24 June 1970, 0906 hours.

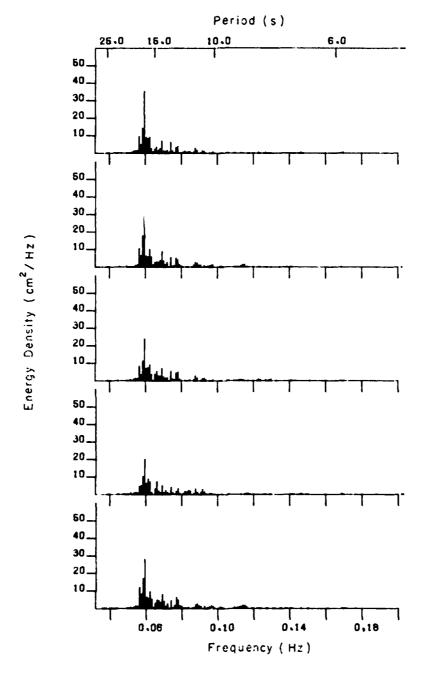


Figure E-18. Significant wave height 69.3 centimeters (2.3 feet), 24 June 1970, 1206 hours.

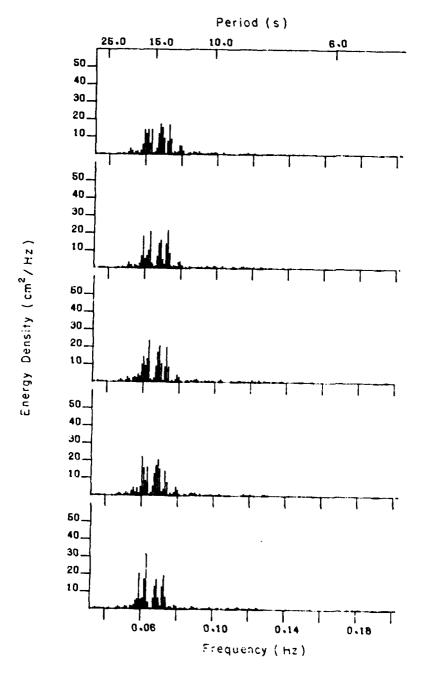


Figure E-19. Significant wave height 79.5 centimeters (2.6 feet), 24 June 1970, 1506 hours.

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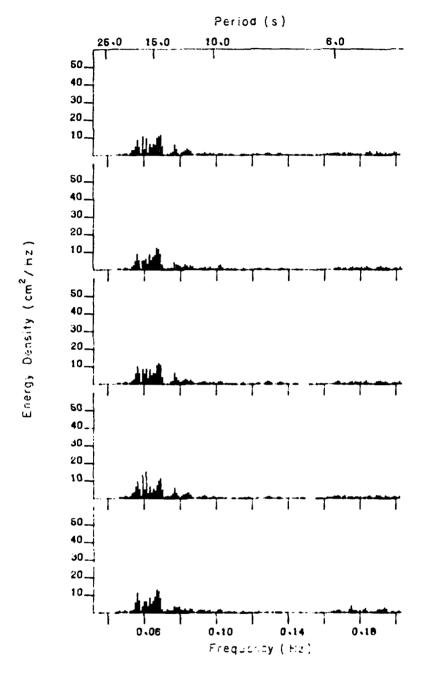


Figure E-20. Significant wave height 72.5 centimeters (2.4 feet), 24 June 1970, 1807 hours.

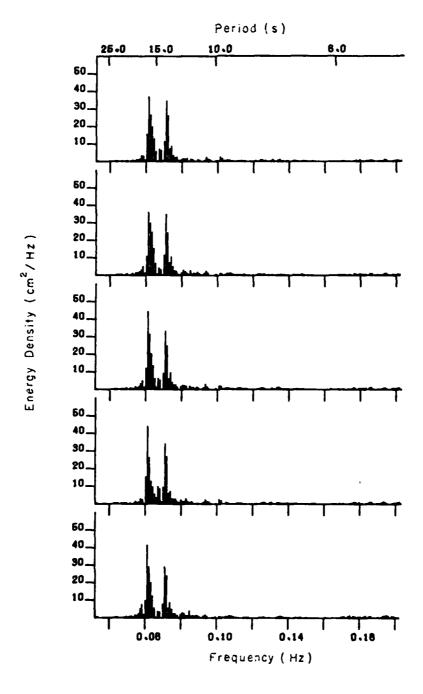


Figure E-21. Significant wave height 70.1 centimeters (2.3 feet), 25 June 1970, 0007 hours.

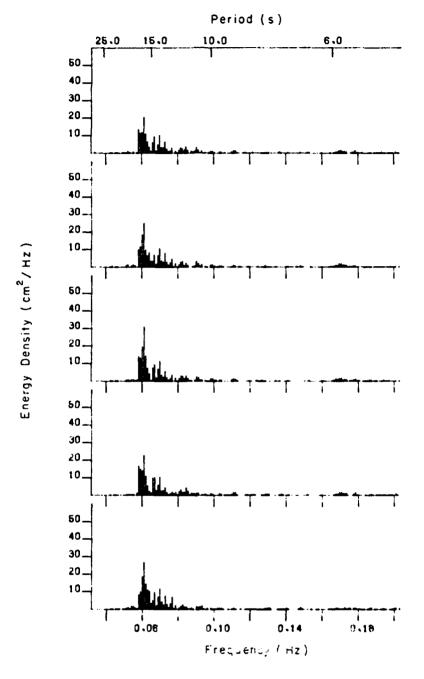


Figure E-22. Significant wave height 76.9 centimeters (2.5 feet), 25 June 1970, 0307 hours.

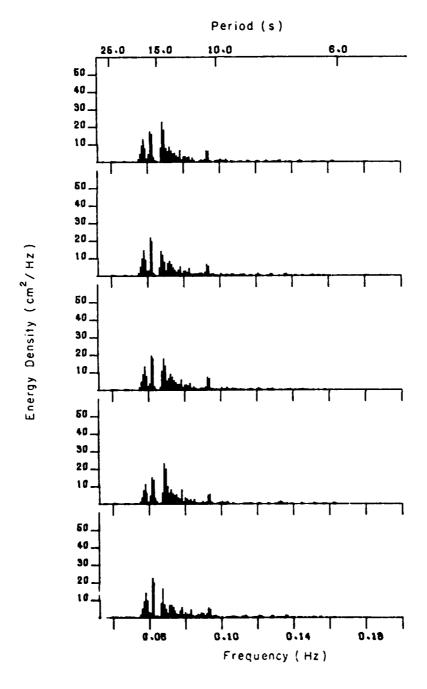


Figure E-23. Significant wave height 66.7 centimeters (2.2 feet), 25 June 1970, 1808 hours.

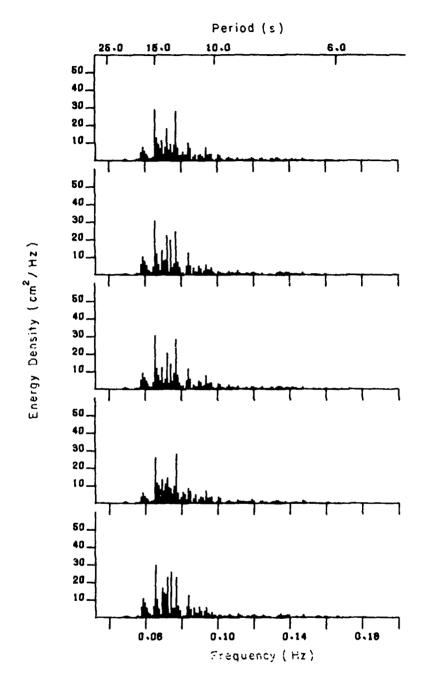


Figure E-24. Significant wave height 70.3 centimeters (2.3 feet), 25 June 1970, 2108 hours.

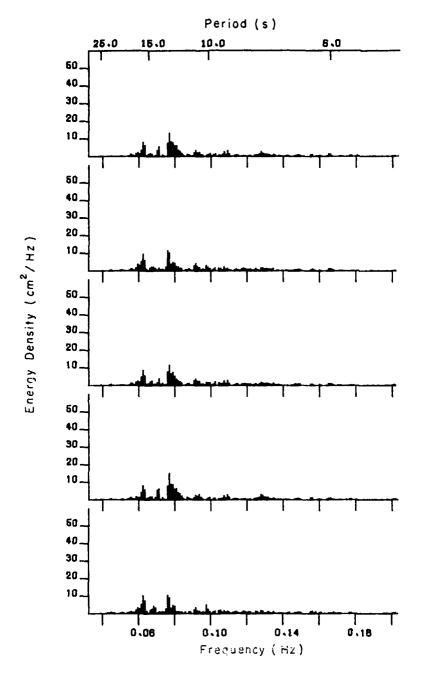


Figure E-25. Significant wave height 61.2 centimeters (2.0 feet), 26 June 1970, 1758 hours.

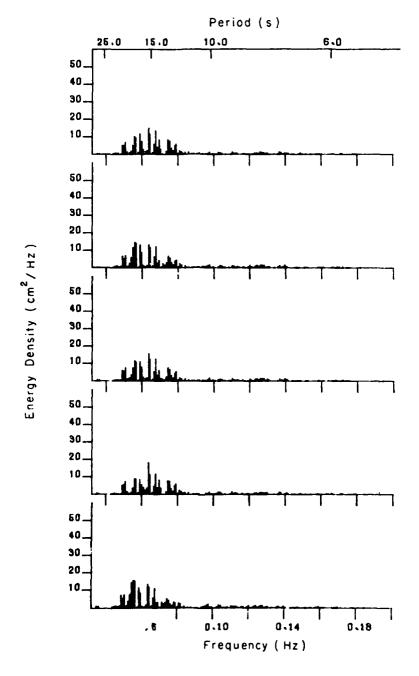
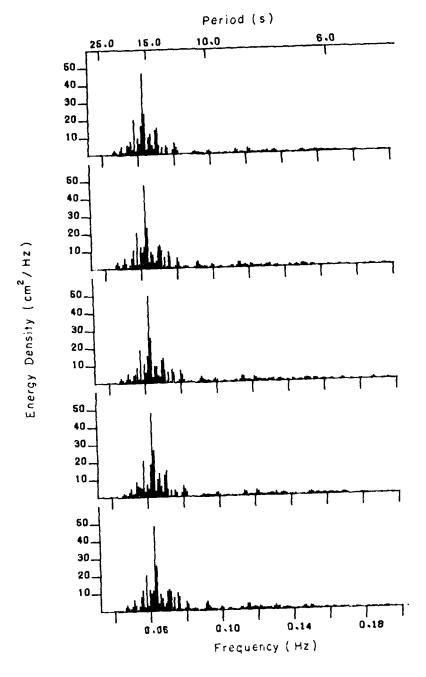


Figure E-26. Significant wave height 79.7 centimeters (2.6 feet), 26 June 1970, 2358 hours.



gure E-27. Significant wave height 92.3 centimeters (3.0 feet), 28 June 1970, 2010 hours.

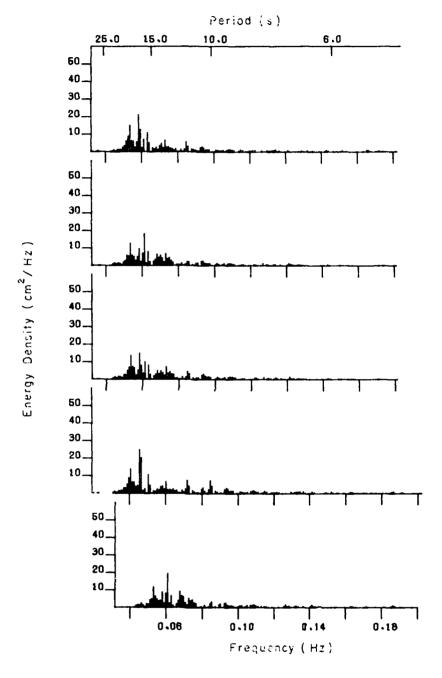


Figure E-28. Significant wave height 83.2 centimeters (2.7 feet), 28 June 1970, 2310 hours.

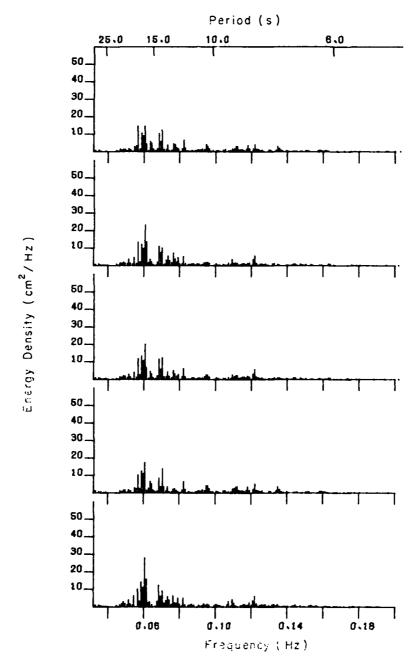


Figure E-29. Significant wave height 78.7 centimeters (2.6 feet), 29 June 1970, 0510 hours.

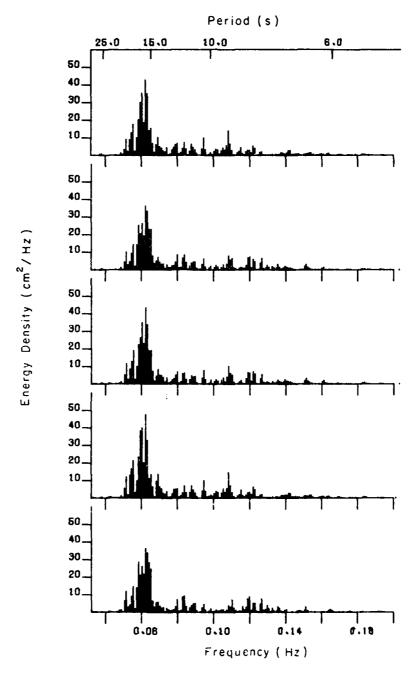


Figure E-30. Significant wave height 64.0 centimeters (2.1 feet), 16 November 1970, 1607 hours.

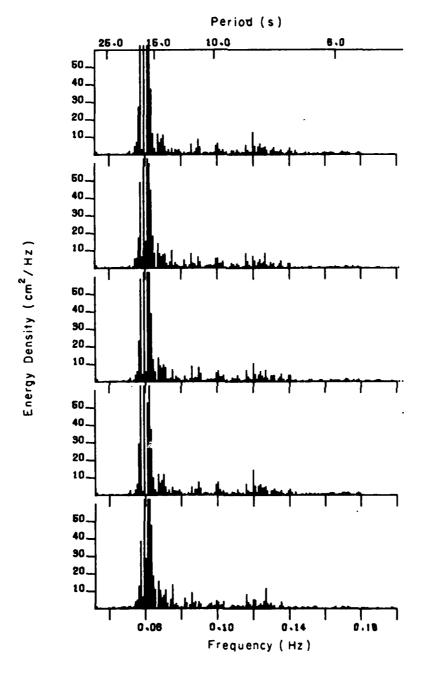


Figure E-31. Significant wave height 80.9 centimeters (2.7 feet), 16 November 1970, 1907 hours.

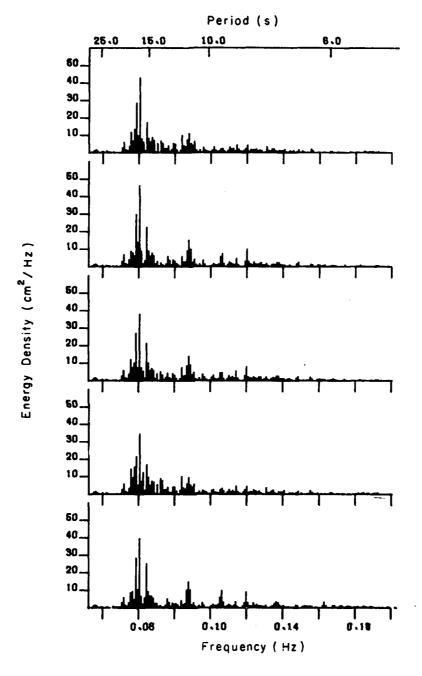


Figure E-32. Significant wave height 63.0 centimeters (2.1 feet), 17 November 1970, 0407 hours.

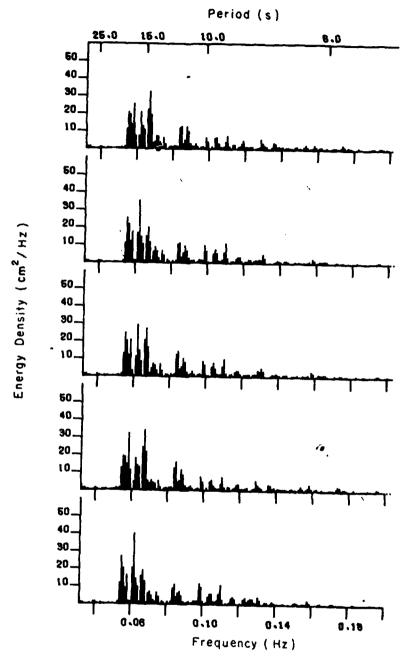


Figure E-33. Significant wave height 74.6 centimeters (2.4 feet), 16 December 1970, 1700 hours.

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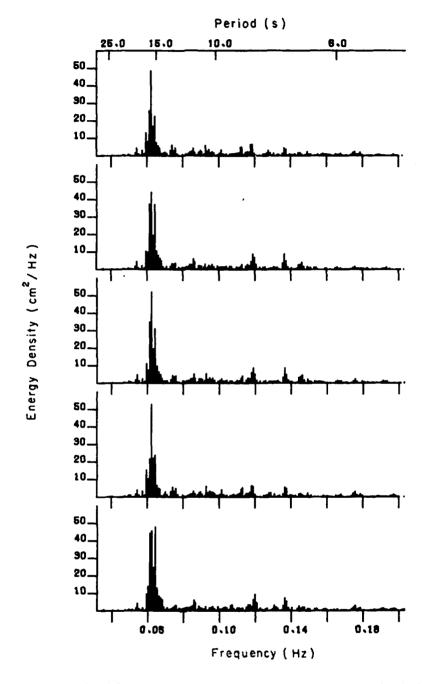


Figure E-34. Significant wave height 95.3 centimeters (3.1 feet), 16 December 1970, 2000 hours.

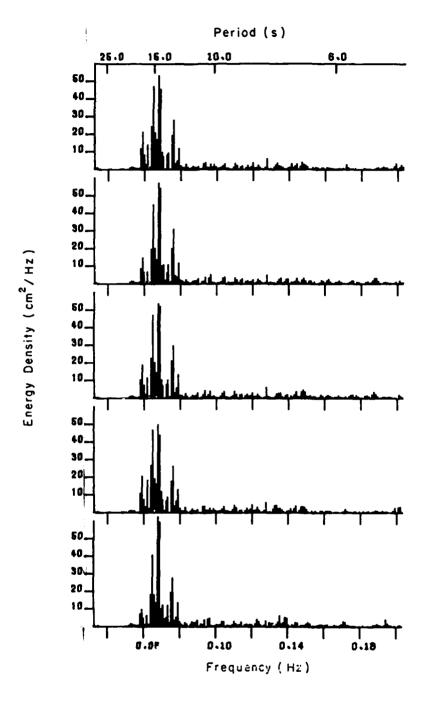


Figure E-35. Significant wave height 85.3 centimeters (2.8 feet), 16 December 1970, 2300 hours.

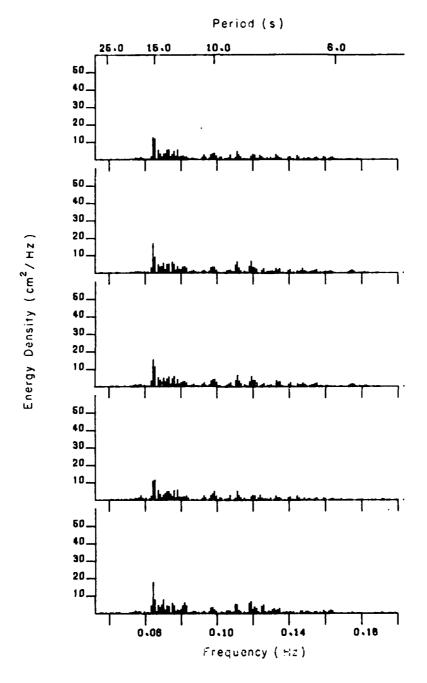


Figure E-36. Significant wave height 90.9 centimeters (3.0 feet), 17 December 1970, 0200 hours.

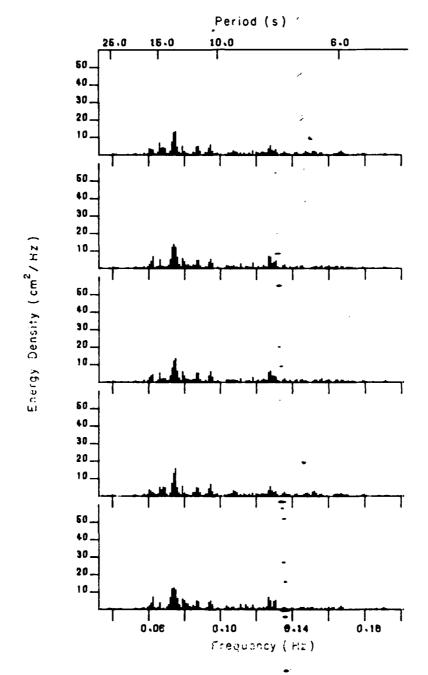


Figure E-37. Significant wave height 103.5 centimeters (3.4 feet), 17 December 1970, 0500 hours.

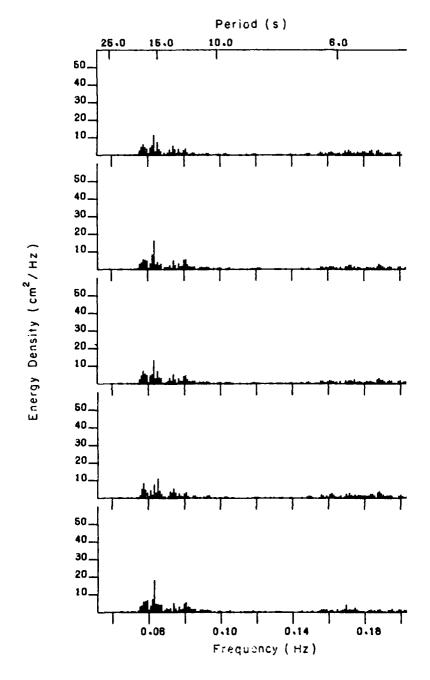


Figure E-38. Significant wave height 120.7 centimeters (4.0 feet), 17 December 1970, 0800 hours.

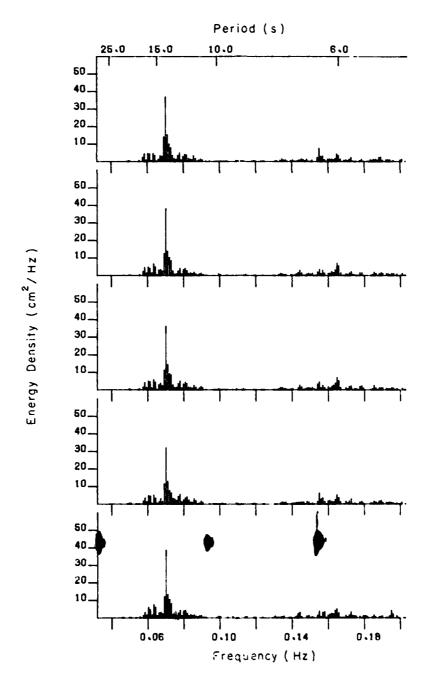


Figure E-39. Significant wave height 97.4 centimeters (3.2 feet), 17 December 1970, 1044 hours.

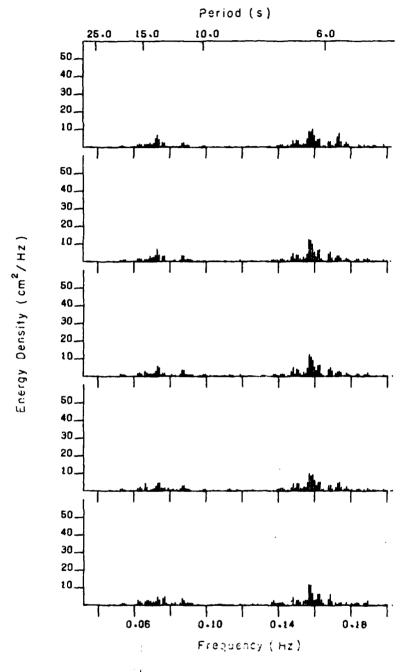


Figure E-40. Significant wave height 93.4 centimeters (3.1 feet), 17 December 1970, 1344 hours.

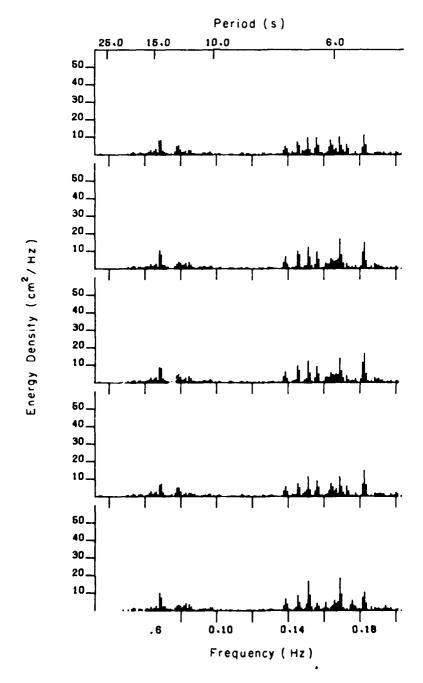


Figure E-41. Significant wave height 91.2 centimeters (3.0 feet), 17 December 1970, 1644 hours.

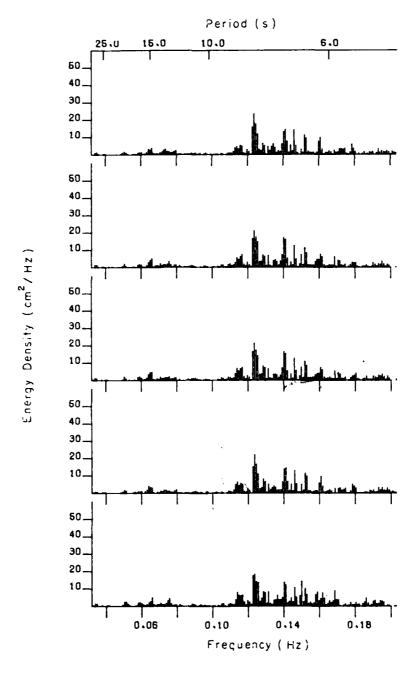


Figure E-42. Significant wave height 108.1 centimeters (3.5 feet), 17 December 1970, 1944 hours.

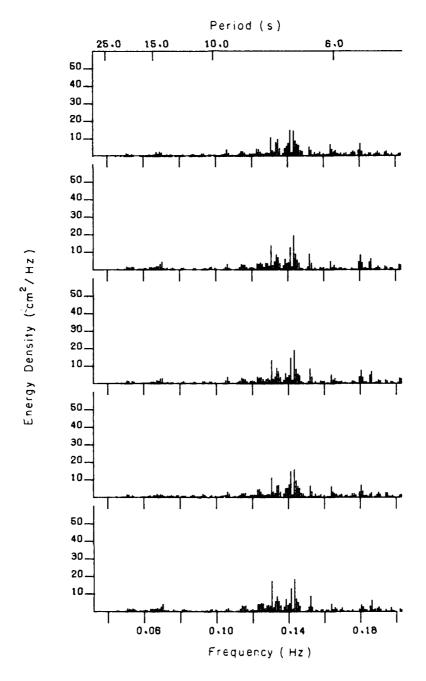


Figure E-43. Significant wave height 71.5 centimeters (2.3 feet), 18 December 1970, 0445 hours.

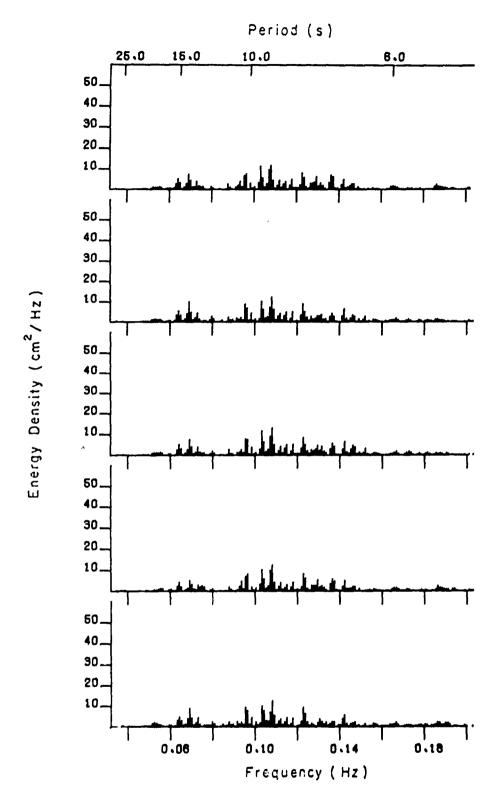


Figure E-44. Significant wave height 66.9 centimeters (2.2 feet), 18 December 1970, 0745 hours.

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TC203 .U581tp no. 77-7 627	Esteva, Dinorah C. Evaluation of the computation of wave direction with three-gage arrays / by Dinorah C. Esteva Fort Belvoir, Va.: U.S. Coastal Engineering Research Center, 1977. 123 p.: ill. (Technical paper - U.S. Coastal Engineering Research Center; no. 77-7) Bibliography: p. 32. A description of the collection and analyses of data obtained with an array of five pressure sensors near Pt. Mugu, California, is presented. The 10 three-gage arrays possible with five gages are used to compare redundant values of the direction of wave propagations. The dependence of directional determination on array orientations. The dependence of directional determination on array orientations. The dependence of directions based on simulated narrox-banded user itside is revealed by calculations based on simulated narrox-banded user treins. 1. Waves. 2. Wave propagation. 3. Wave direction. 4. Pt. Mugu, Calif. 1. Title. 11. Series: U.S. Goastal Engineering Research Center. Technical paper no. 77-7.	TC203 .U581tp no. 77-7 627
TC203 .U531tp no. 77-7 627	Esteva, Dinorah C. Evaluation of the computation of wave direction with three-gage arrays / by Dinorah C. Esteva Fort Belvoir, Va.: U.S. Coastal Engineering Research Center, 1977. 123 p.: 111. (Technical paper - U.S. Coastal Engineering Research Center; no. 77-7) Bibliography: p. 32. A description of the collection and analyses of data obtained with an array of five pressure sensors near Pt. Mugu, California, is presented. The 10 three-gage arrays possible with five gages are used to compare redundant values of the direction of wave propagations. The dependence of directional determination on array orientation relative to incident wave direction and wavelength at the array site is revealed by calculations based on simulated narrow-banded wave trains. I. Waves. 2. Wave propagation. 3. Wave direction. 4. Pt. Mugu, Calif. I. Hille. II. Series: U.S. Coastal Engineering Research Center. Technical paper no. 77-7.	TC203 .U531tp no. 77-7 627
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